

~~PROSPECTIVE~~ NASA CR 137762

Final Report

MEASUREMENTS OF TRACE CONSTITUENTS
FROM ATMOSPHERIC INFRARED
EMISSION AND ABSORPTION SPECTRA -
A FEASIBILITY STUDY

A. Goldman, W.J. Williams and D.G. Murcay

September 1974

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Prepared under Contract No. NAS2-8200

by _____

Department of Physics and Astronomy
University of Denver
Denver, Colorado 80210

for

AMES RESEARCH CENTER
NATIONAL AERONAUTICS AND SPACE ADMINISTRATION

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Measurements of Trace Constituents from Atmospheric Infrared
Emission and Absorption Spectra - A Feasibility Study

by

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Abstract

The feasibility of detecting eight trace constituents (CH_4 , HCl , HF , HNO_3 , NH_3 , NO , NO_2 and SO_2) against the rest of the atmospheric background at various altitudes from infrared emission and absorption atmospheric spectra has been studied. Line-by-line calculations and observational data have been used to establish features that can be observed in the atmospheric spectrum due to each trace constituent. Model calculations have been made for experimental conditions which approximately represent "state of the art" emission and absorption spectrometers.

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I. Introduction

Infrared methods are very powerful techniques for trace gas measurement due to the unique patterns in the fine structure of the spectral lines of both the background atmosphere and trace constituents. The enormous amount of work of many individuals during the last decade has resulted in extensive tabulations of spectral line parameters of a number of gases. During that time, efficient line-by-line computer programs have been developed which, with today's large computers, allow the generation of synthetic atmospheric spectra for arbitrary atmospheric profiles of minor constituents with arbitrary optical paths and spectral resolutions. While these techniques have been extensively used for the analysis of observed atmospheric emission and absorption infrared spectra, their use in feasibility studies for infrared measurements of atmospheric trace constituents⁽¹⁻¹⁸⁾ was quite limited.

When small amounts of gas are to be measured, accurate knowledge is required of the fine details of both the experiment and the theory. In particular, the effects of interference between various molecular features that might be present in the spectrum require special attention. The line-by-line techniques are very powerful in such cases, as is demonstrated in a recent study of the solar spectrum during a search for stratospheric NO.⁽¹⁹⁾ It was shown that a detailed line-by-line analysis of interferences is a basic requirement for a meaningful determination of minor constituents from weak spectral features.

A similar problem occurs with many other gases in various spectral regions. Therefore, in the present work we have applied the line-by-line techniques to the problem of finding detectable limits for a number of trace constituents against the rest of the atmospheric background. In particular, we have studied the effects of interference on the detectability of small amounts of CH₄, HCl, HF, HNO₃, NH₃, NO, NO₂ and SO₂ under

conditions corresponding to selected practical atmospheric experiments with absorption and emission spectrometers.

II. Calculation Procedure

In the present calculations a one-layer model of the atmosphere was used with appropriate effective temperature, pressure and mixing ratio to determine the optical path. The effective mixing ratios for the minor constituents of the background atmosphere were derived from the AFCRL mid-latitude summer profile.⁽²⁰⁾

The spectra were calculated on a line-by-line basis using the most recently available line parameters for the various molecular species. The line parameters for the background atmosphere are those generated by various researchers (notably by W.S. Benedict) and compiled on magnetic tape by AFCRL.⁽²¹⁾ These cover most of the relevant bands of H₂O, CO₂, O₃, N₂O, CH₄, CO and O₂. Additional line parameters that were generated at the University of Denver and elsewhere for the strongest bands of HCl,⁽²²⁾ HF,⁽²³⁾ NH₃,⁽²⁴⁾ NO,⁽²⁵⁾ NO₂,⁽²⁶⁾ and SO₂^(27,28) were used for the calculations of the spectra of the background atmosphere with the additional constituents. A Voigt line shape was used for all spectral lines.

It should be noted that the calculated background spectra presented here do not include the HNO₃ vibration-rotation bands, the H₂O vapor continuum and the pressure-induced bands of O₂ and N₂. Deviations from the Voigt line shape are also neglected here. In addition, solar molecular and atomic lines, which can cause significant interference with atmospheric lines in actual atmospheric spectra taken with the sun as a source, have not been included. The corresponding NO and HF spectra presented here should have been thus superimposed on the solar CO $\Delta v=1$ and $\Delta v=2$ transitions respectively.

The effects of interference between solar CO lines and atmospheric NO lines has been studied earlier.⁽¹⁹⁾ In addition, numerous atomic solar lines occur in the vicinity of the HF and HCl lines presented here. The effects of these lines on the presently used HF and HCl lines are discussed here only briefly because the subject requires a separate detailed study. Furthermore, it should be noted that for a constituent which is layered sharply at a certain altitude, the constant mixing ratio assumed for the trace constituents added to the background atmosphere makes the present charts less accurate.

The calculations were performed so as to correspond to three different experimental techniques; long path absorption cell with laser spectroscopy, atmospheric absorption spectra using the sun as a source and atmospheric emission spectra. The calculations were performed for "infinite" resolution and the resolution was then degraded to correspond to resolutions of 0.08 cm^{-1} and 1.0 cm^{-1} , which are close to current "state of the art" of each experimental technique. Examples of each technique are found in the literature.^(1-19, 29-39) The absorption data were combined with the appropriate blackbody function to get the atmospheric emission data. Four altitudes were chosen as the base of each atmospheric layer: 5.5 km, 15 km, 25 km and 35 km. These were chosen to represent the upper troposphere, near or just above the tropopause, the ozone layer and the middle stratosphere where photochemical effects dominate. The main advantage of solar spectroscopy and atmospheric emission spectra lies in the large optical path which is achieved at large zenith angles. These calculations were therefore performed for zenith angles at 80° and 85° . For laser spectroscopy with a long path cell, a cell of 100 meters was chosen for all cases except for CH₄ where shorter cells were chosen.

III. Results

The theoretical calculations are presented in graphical form showing either radiance or emissivity as a function of wavelength and wavenumber. The calculations are presented according to molecule, altitude and the three experimental techniques discussed above. In most cases the data are plotted on both logarithmic and linear scales to demonstrate graphically how the limit of detectability depends on the dynamic range which can be achieved with the observational technique employed. In each case the atmospheric background spectrum is plotted on the same figure with the spectrum of background plus the trace constituent. The stronger features of the particular trace constituent are denoted by vertical or horizontal arrows. The vertical scales of these plots are offset for clarification. In addition, the experimental conditions are denoted on each figure and include the spectral resolution, the altitude and the optical path.

When either logarithmic or linear plots are shown on a chart, the scales on the left and on the right sides apply to the atmosphere with and without the trace constituent respectively. When both logarithmic and linear plots are shown on the same chart, the log and the linear scales are on the left and the right sides respectively. In these cases, the upper and the lower scales apply to the atmosphere with and without the trace constituent respectively.

The spectral calculations for each trace constituent and the associated specific discussions pertaining to each species are found below in separate sections. Detectable amounts of the individual molecules are discussed on the basis of the atmospheric emission and absorption spectra calculated in Figures 1 through 34. In preparing these figures the criterion for a measurable absorptivity was taken as ~1%. The measurable radiance level depends on the wavelength region and is discussed for each molecule separately.

III. A. CH₄

The natural abundance of CH₄ in the lower stratosphere is of the order of 0.1 to 1.0 ppmv. Such a mixing ratio is larger than the mixing ratios of all the other molecules considered in the present report, as the others are trace gases to be measured against the rest of the atmospheric gases. As a result, a relatively short path length and medium resolution can be used for atmospheric CH₄ measurements, either from the solar spectrum or from the atmospheric emission spectrum. A large number of such spectra have been observed from various altitudes, with the most intense CH₄ features occurring in the 3.3 μ m and the 7.6 μ m regions. Therefore in the present report these two regions are considered only for laser spectroscopy with an absorption cell. The calculated CH₄ absorptions presented here are weak and allow linear scaling of gas amount and absorption up to ~10% absorption.

"Infinite" resolution atmospheric absorption calculations for a 5 meter path for both the 3.3 μ m and 7.6 μ m Q-branch regions, at 5.5 km altitude, are shown in Figures 1 and 2 and Figure 3 respectively. Assuming ~1% measurable absorptivity allows the measurement of ~0.5 ppmv CH₄ under these conditions. It is seen that both CH₄ regions are subject to significant interference from atmospheric lines, but the interference is less in the 3.3 μ m region.

Similar "infinite" resolution calculations at 25 km altitude for a 40 meter path near 3.3 μ m and a 20 meter path near 7.6 μ m are shown in Figures 4 and 5 and Figure 6 respectively. Scaling these figures to ~1% measurable absorptivity allows the measurement of ~0.5 ppmv CH₄ with an ~20 meter cell from either region. It is noted that while the atmospheric interference is now practically negligible for the 3.3 μ m region (where the numerous atmospheric lines are lower in absorption intensity by orders of magnitude in comparison with the CH₄ lines),

the $7.6\mu\text{m}$ CH_4 lines are superimposed on many atmospheric lines with absorption intensity comparable to that of the CH_4 lines.

The line parameters for CH_4 and the atmosphere are relatively accurate in both spectral regions. It should be noted, however, that the Q-branch CH_4 lines are subject to uncertainties, but their accumulated intensity permits a lower detectable limit than that for P- and R-branch lines.

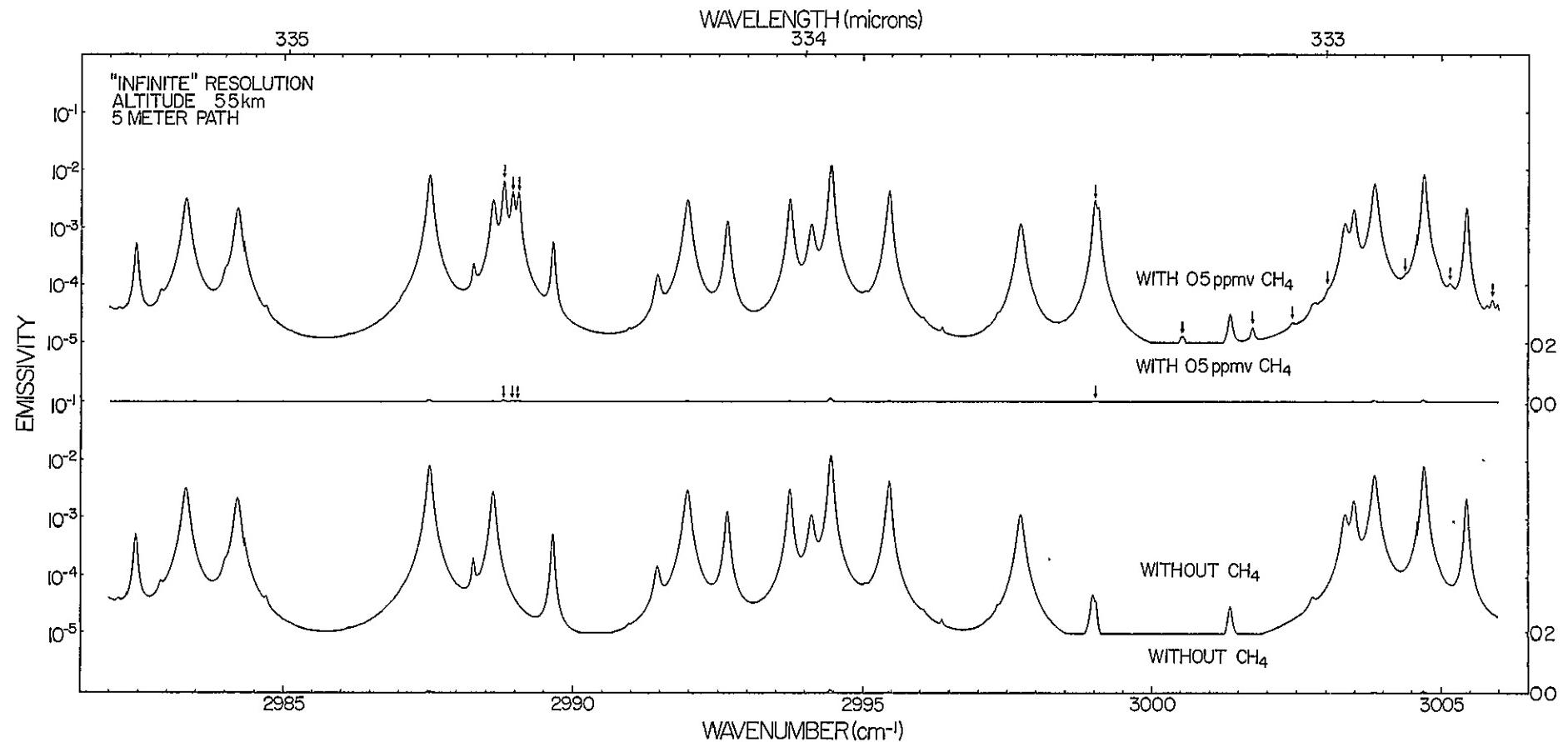


Figure 1.

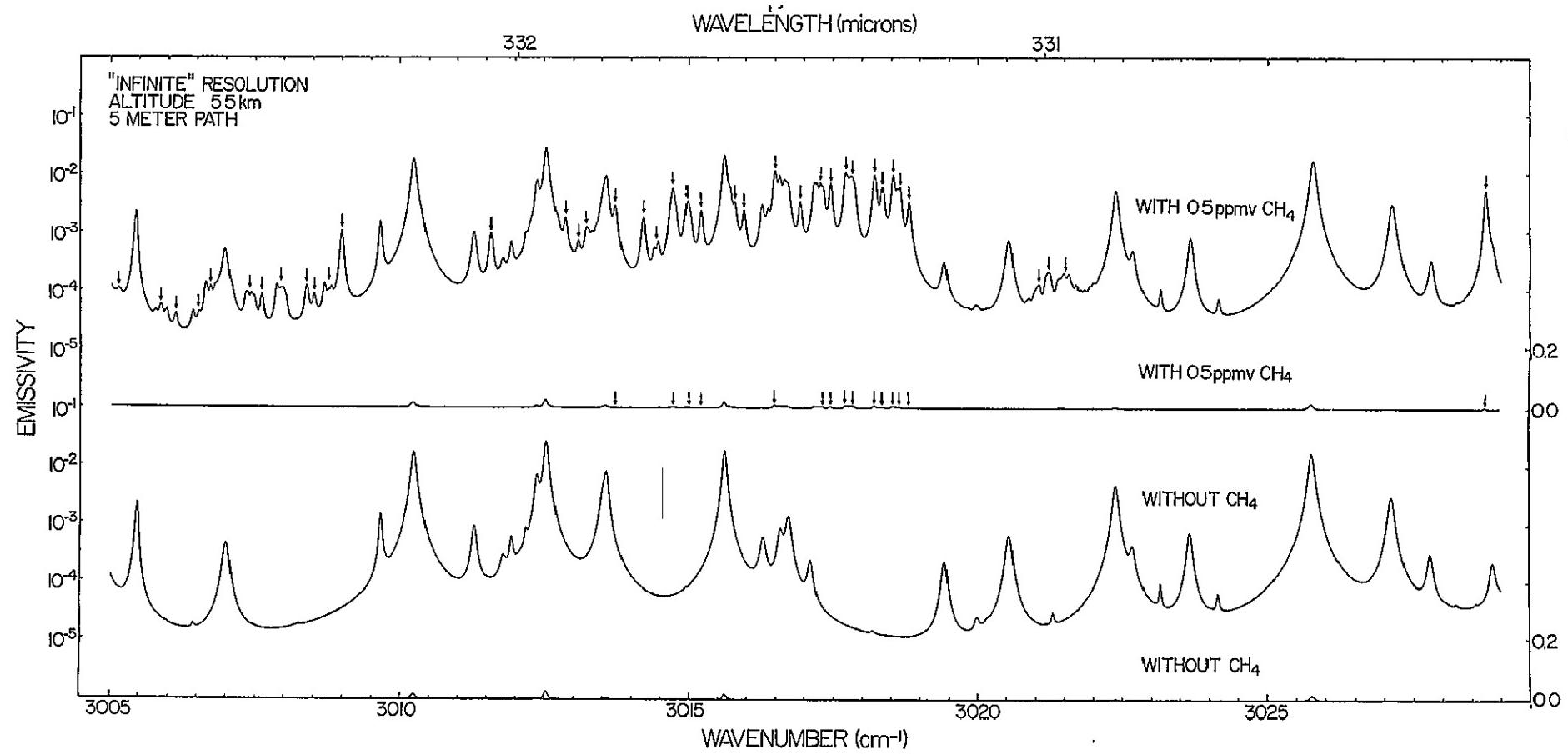


Figure 2.

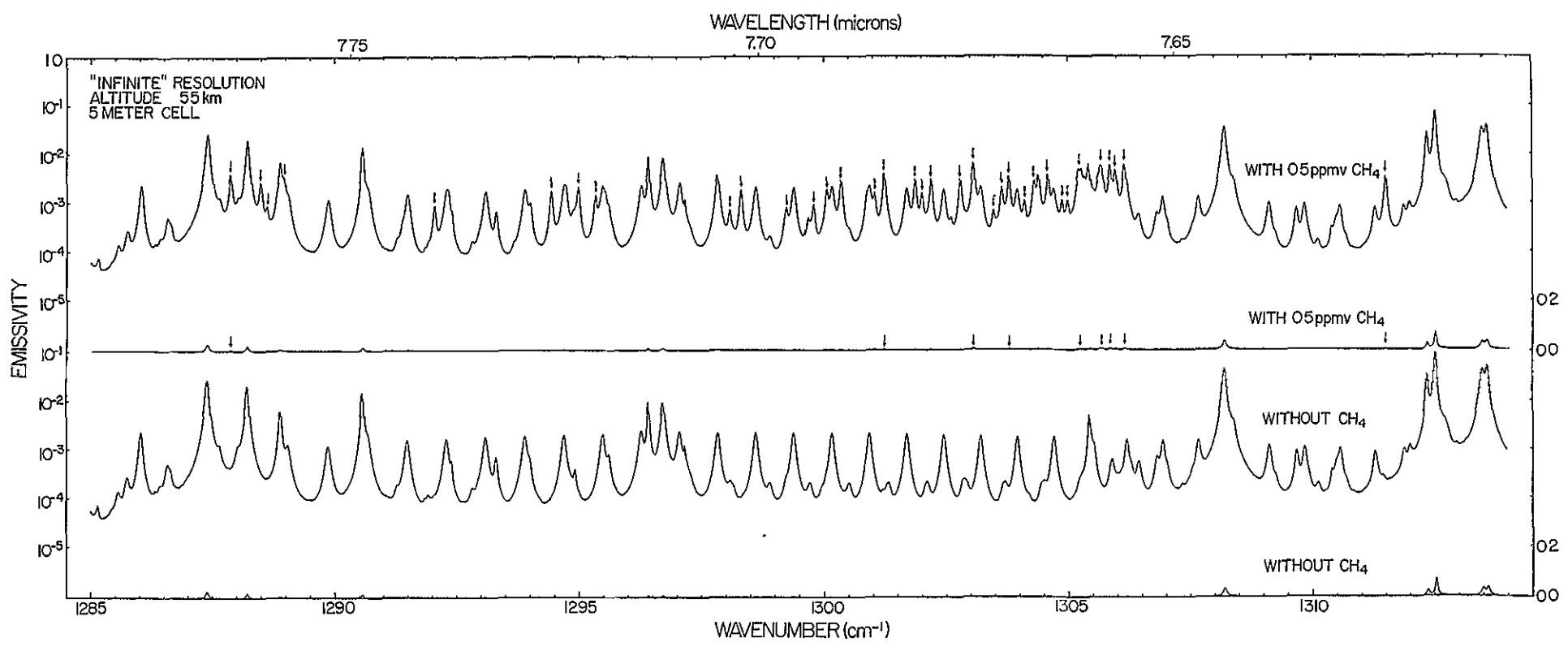


Figure 3.

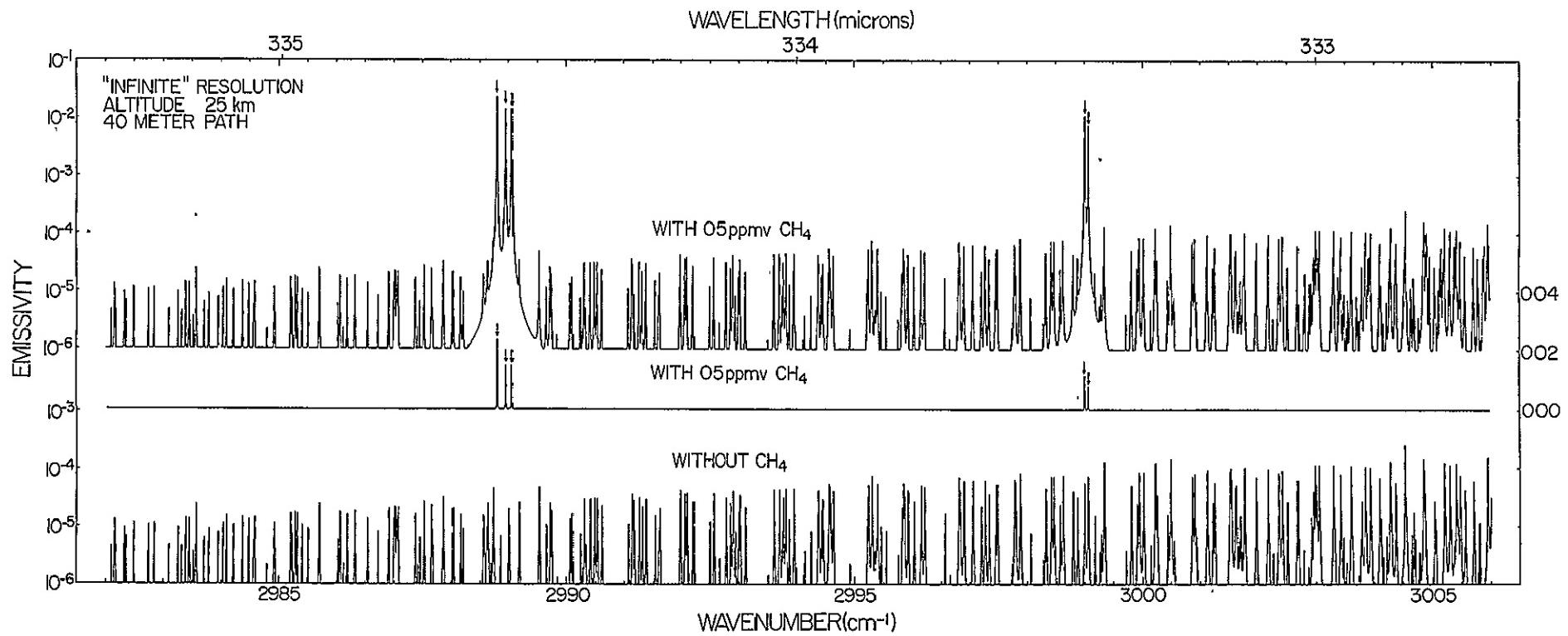


Figure 4.

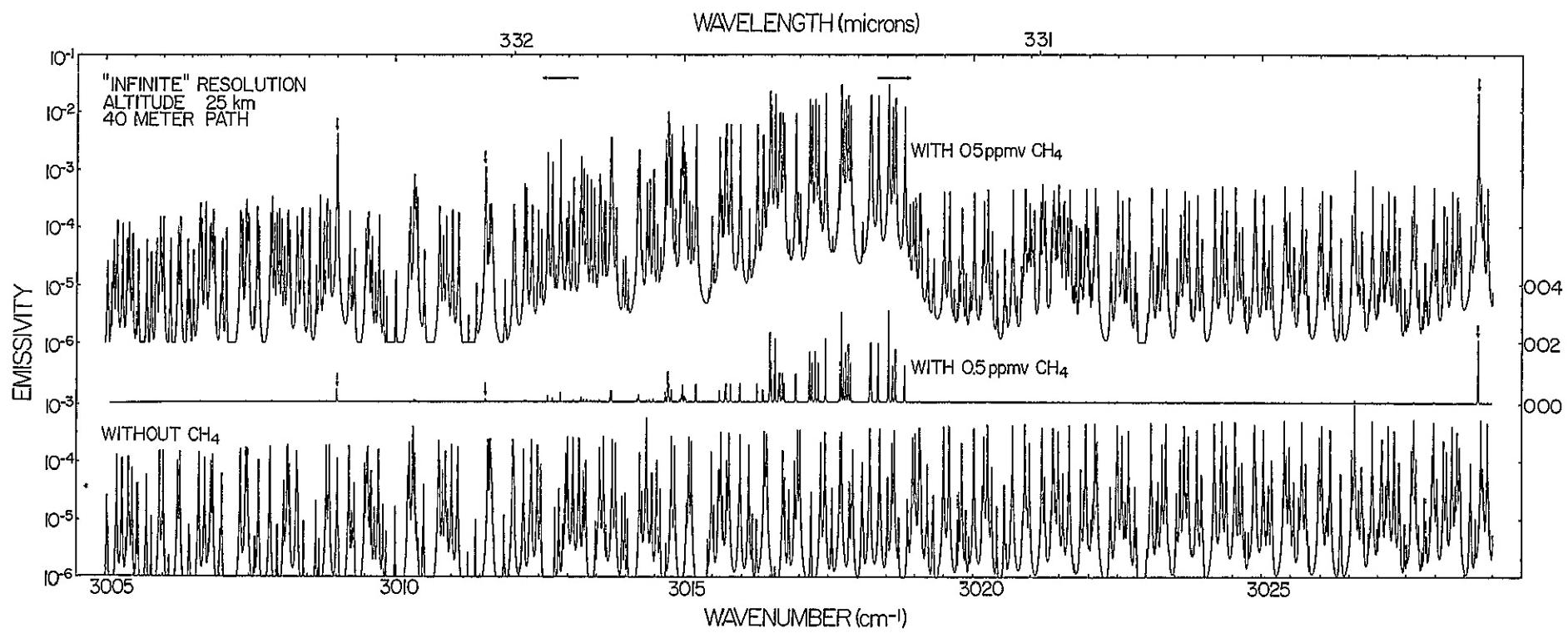


Figure 5.

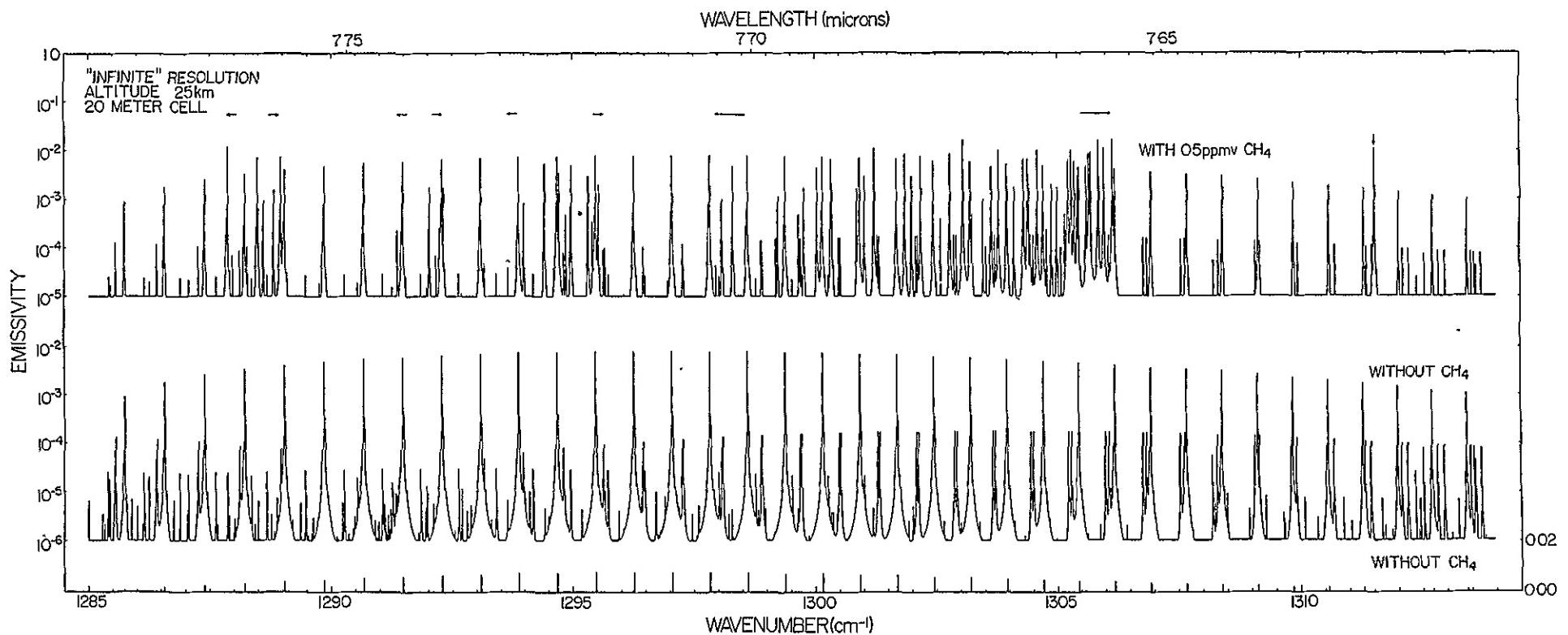


Figure 6.

III. B. HCl

The (0-1) vibration-rotation band of HCl offers several possibilities for HCl detection.

"Infinite" resolution atmospheric absorption calculations for a 100 meter path at 5.5 km and 15 km altitudes are shown in Figures 7 through 9. Assuming an ~1% measurable absorptivity allows the measurement of ~10 ppbv and ~20 ppbv HCl at 5.5 and 15 km by scaling these figures to ~1% emissivity. It is seen that in the spectral region under consideration there is practically no interference from atmospheric lines for a 100 meter path.

Atmospheric spectral radiance calculated with and without HCl for 15 km altitude, 80° zenith angle and 1 cm^{-1} resolution is shown in Figures 10 and 11. It is seen that 1 ppbv of HCl gives rise to a number of HCl features superimposed on the atmospheric spectrum. The radiance levels shown for the HCl features at $\sim 2 \times 10^{-9} \text{ w cm}^{-2} \text{ sr}^{-1} \mu\text{m}^{-1}$ are at the lower limit of "state of the art" spectral radiometers at $3.5 \mu\text{m}$.

By dividing the radiance signal by the blackbody signal, the above radiance figures yield an absorptivity of ~1%. Therefore, a solar spectrum taken under these conditions (15 km altitude, 80° zenith angle and 1 cm^{-1} resolution) will allow the measurement of ~1 ppbv HCl. Such spectra are easily obtained with current solar spectrometers.

Absorption spectra as would be expected in solar spectrum measurements from 25 km at 85° zenith angle and 0.08 cm^{-1} resolution, are calculated in Figures 12 through 14 for two spectral regions. From Figures 12 and 13 it is seen that the HCl isotopic lines near 2945 cm^{-1} are well separated from the atmospheric lines. Absorptivity of ~1% allows the measurement of ~0.1 ppbv HCl. A comparison of Figures 12 and 13 and Figure 14 shows that the 2945 cm^{-1} region is more useful than the 2820 cm^{-1} region for HCl measurements in the atmosphere.

While the intensities of the HCl lines in the two spectral regions are quite similar, the interference from the atmospheric lines is significantly smaller in the 2945 cm^{-1} region.

The line parameters for HCl and the atmosphere are relatively accurate in the $3.4\mu\text{m}$ region. However, closer attention should be given to the solar background lines for solar spectra measurements in the HCl region where several atomic lines have been observed.⁽³⁹⁾

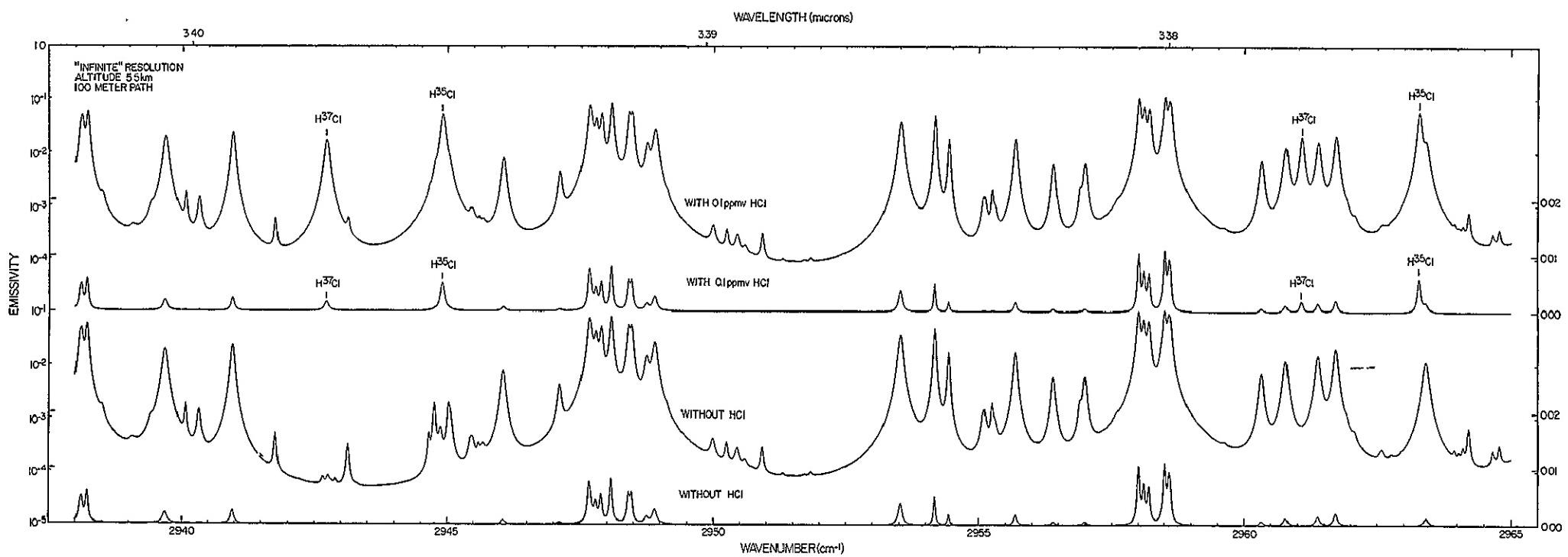


Figure 7.

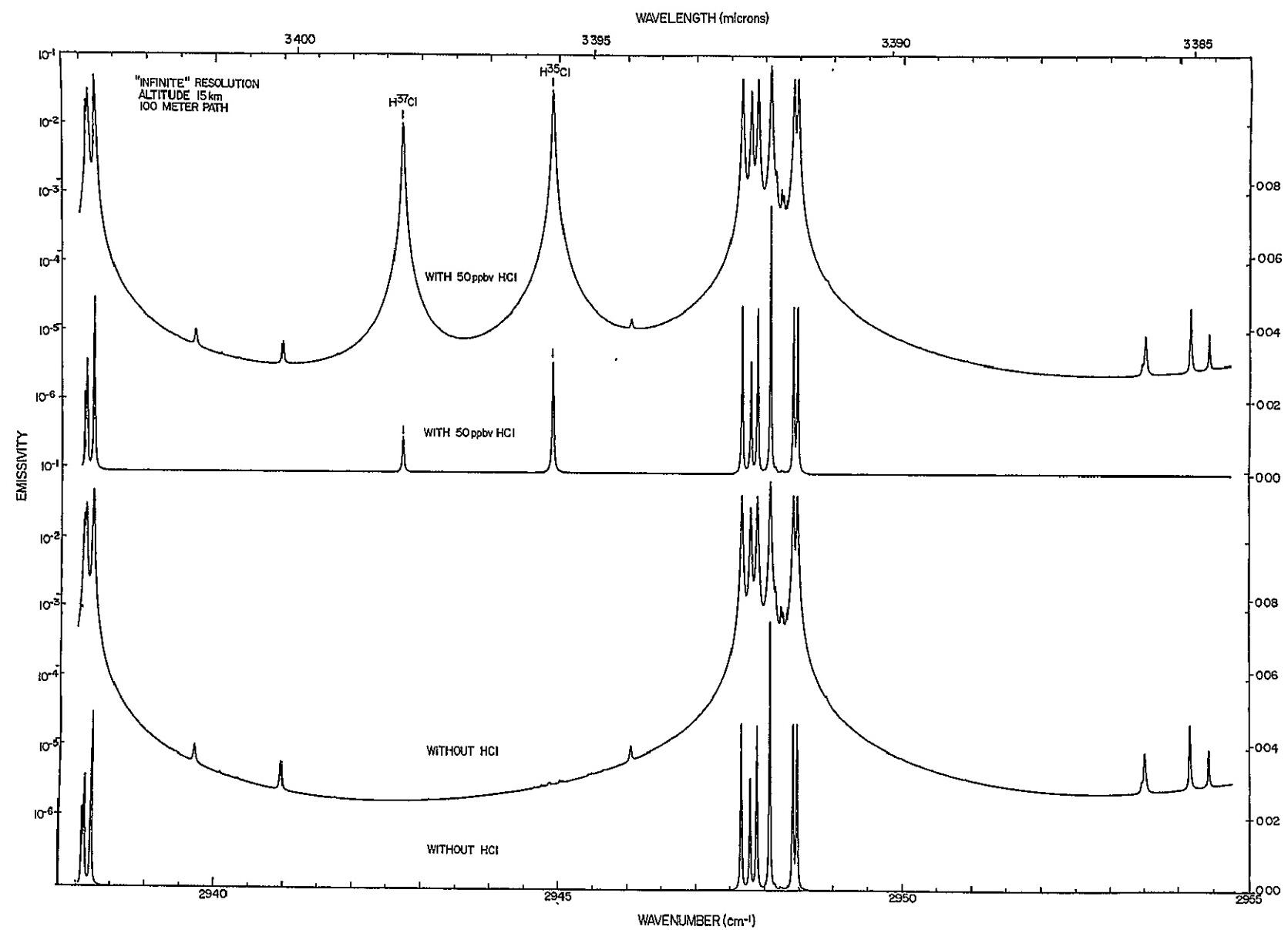


Figure 8.

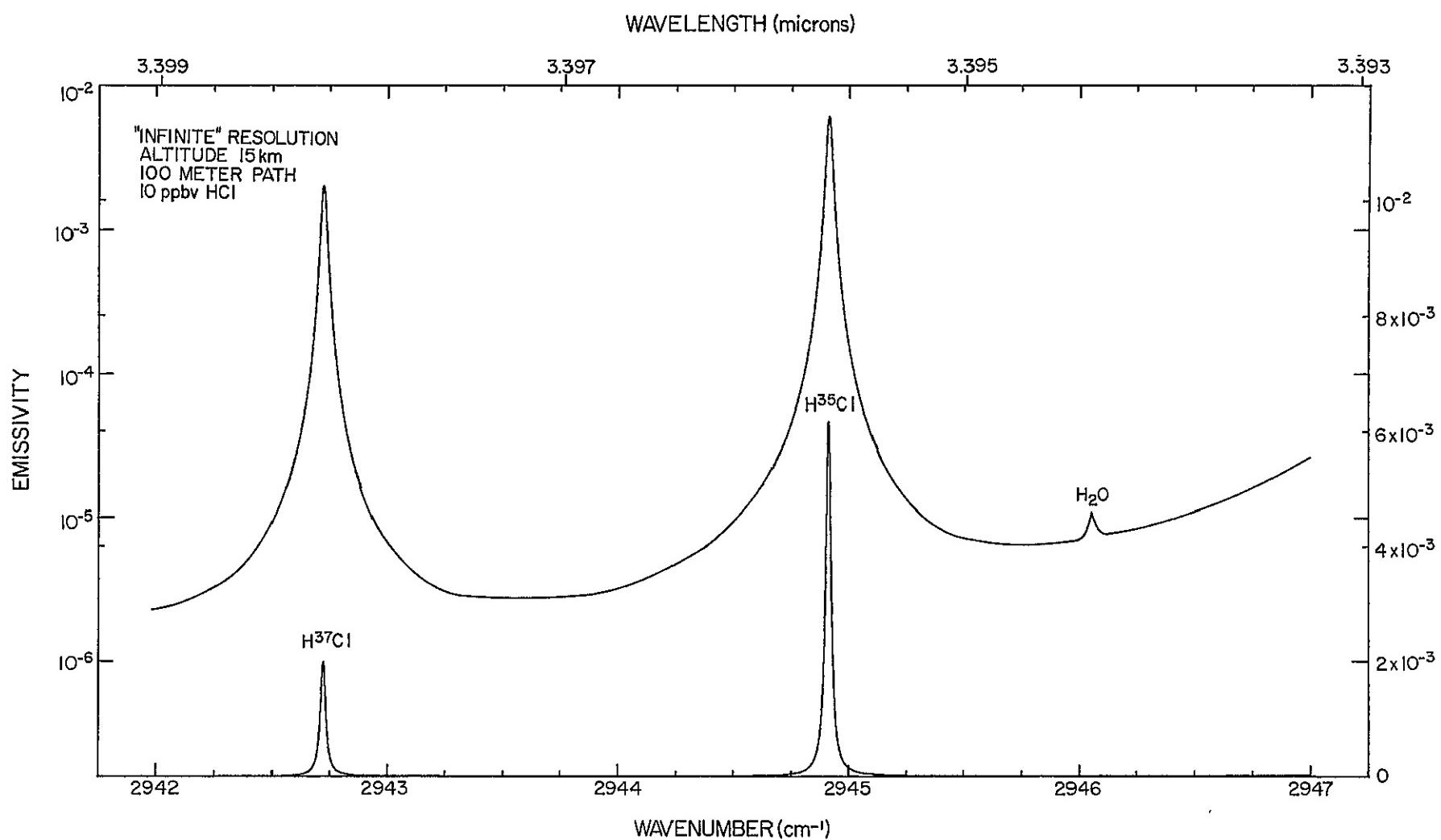


Figure 9.

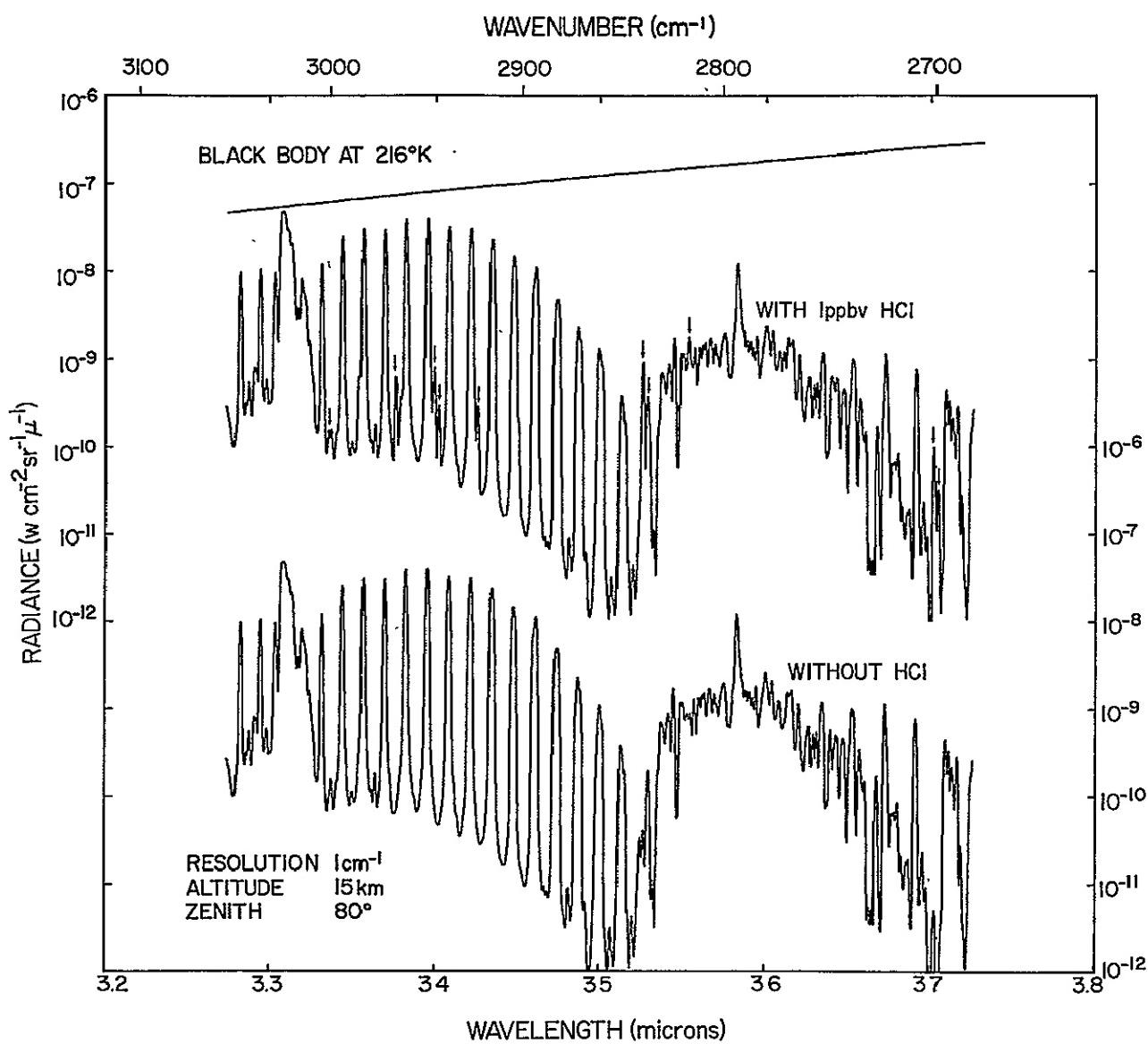


Figure 10.

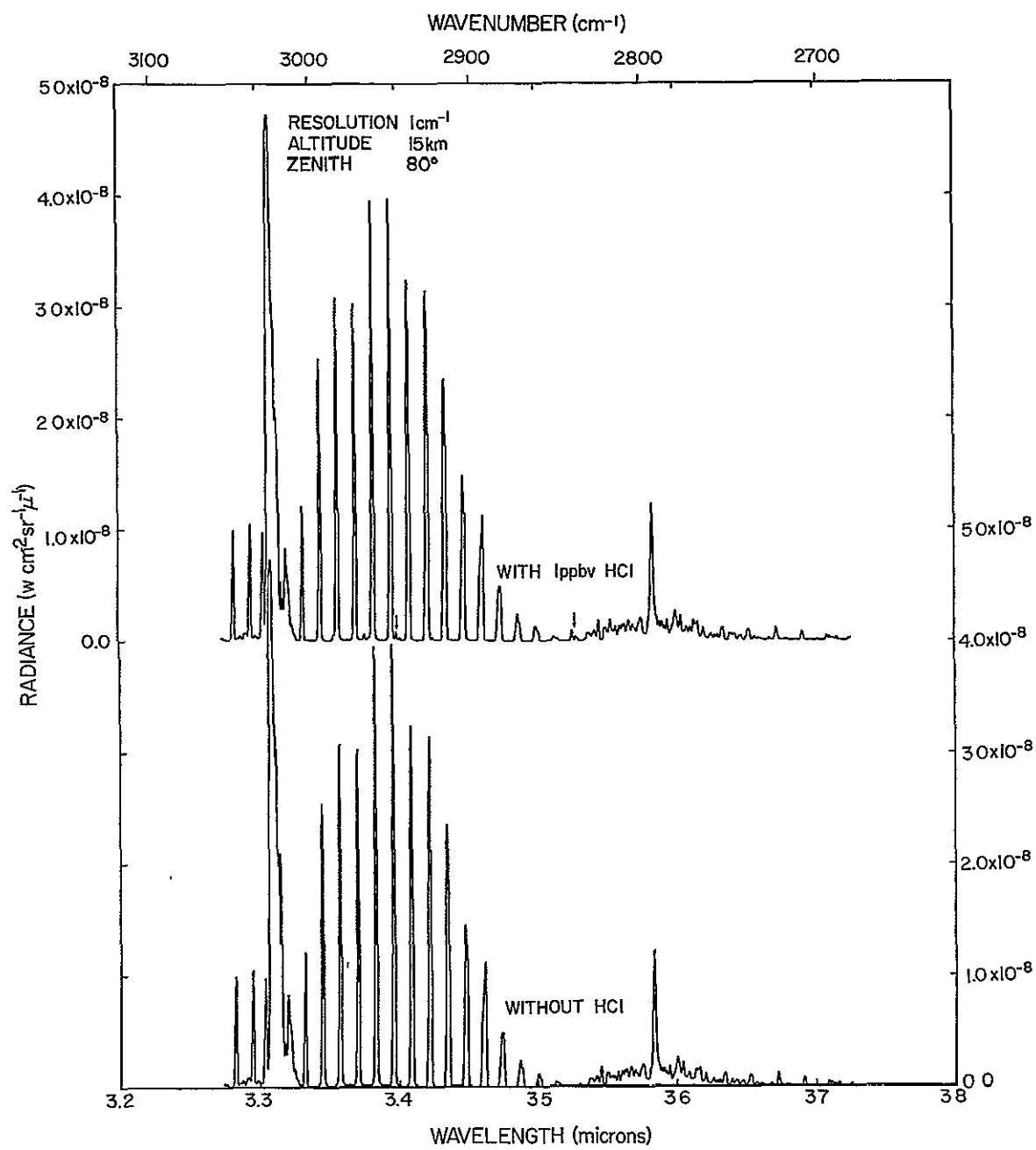


Figure 11.

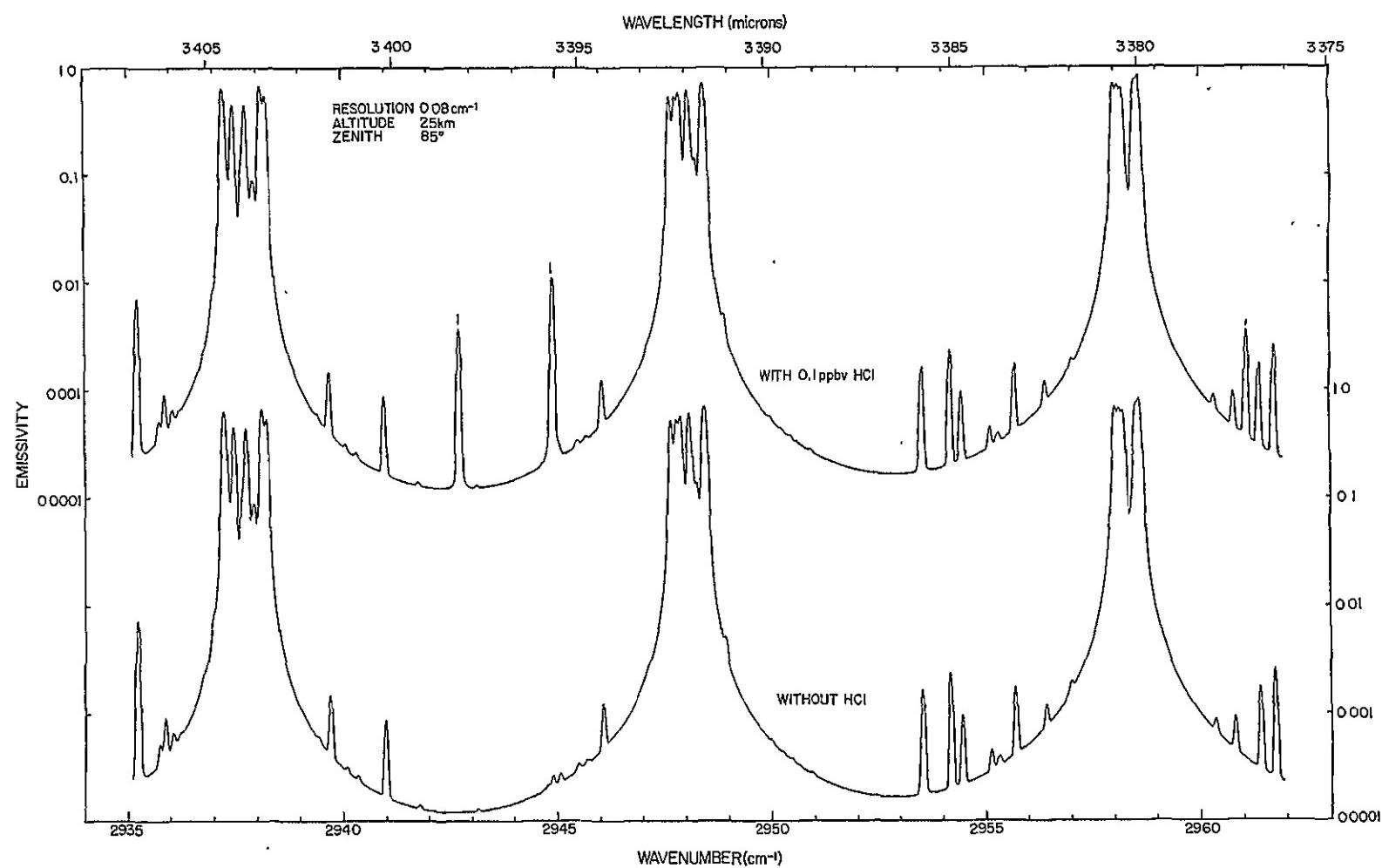


Figure 12.

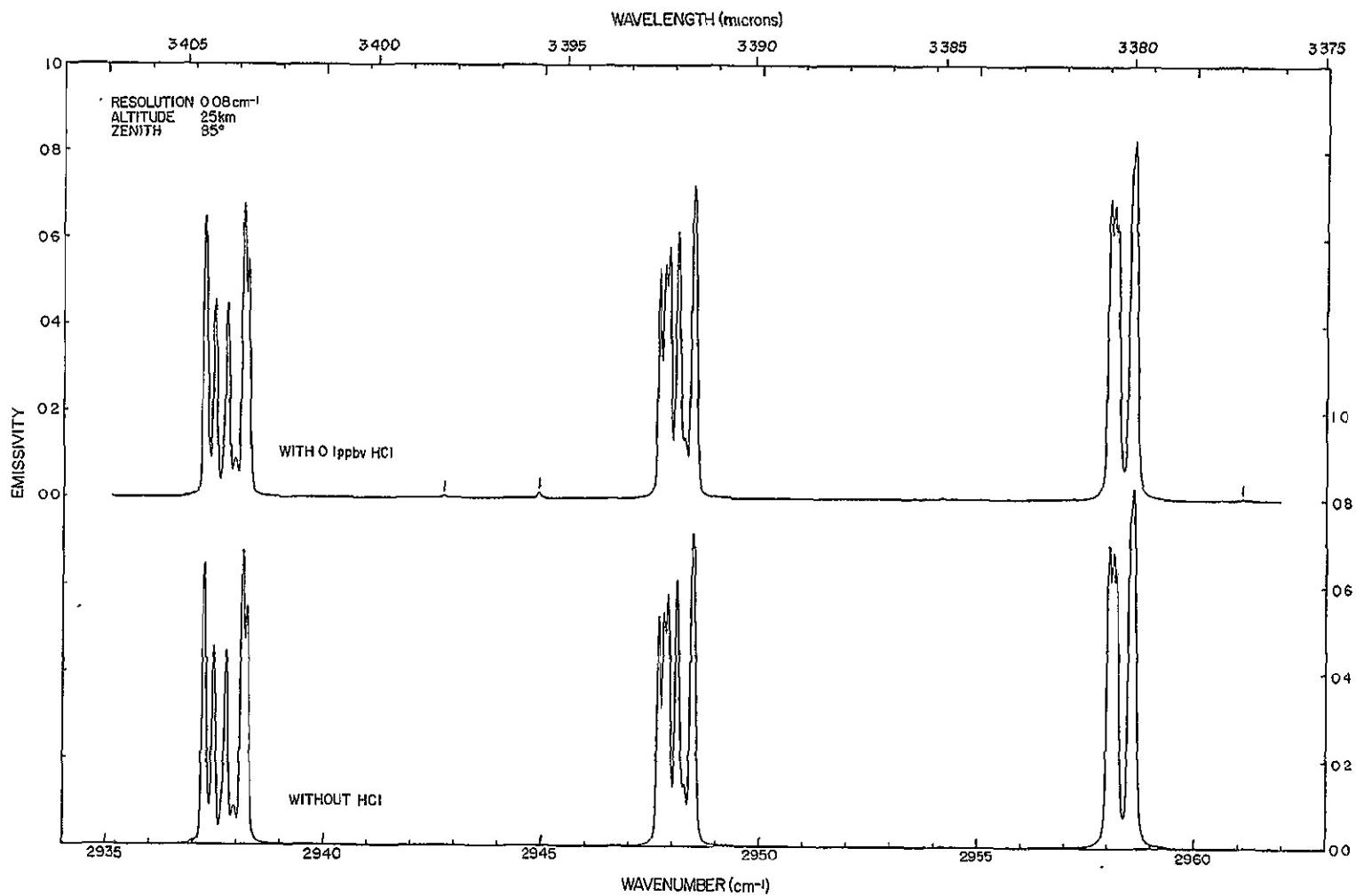


Figure 13.

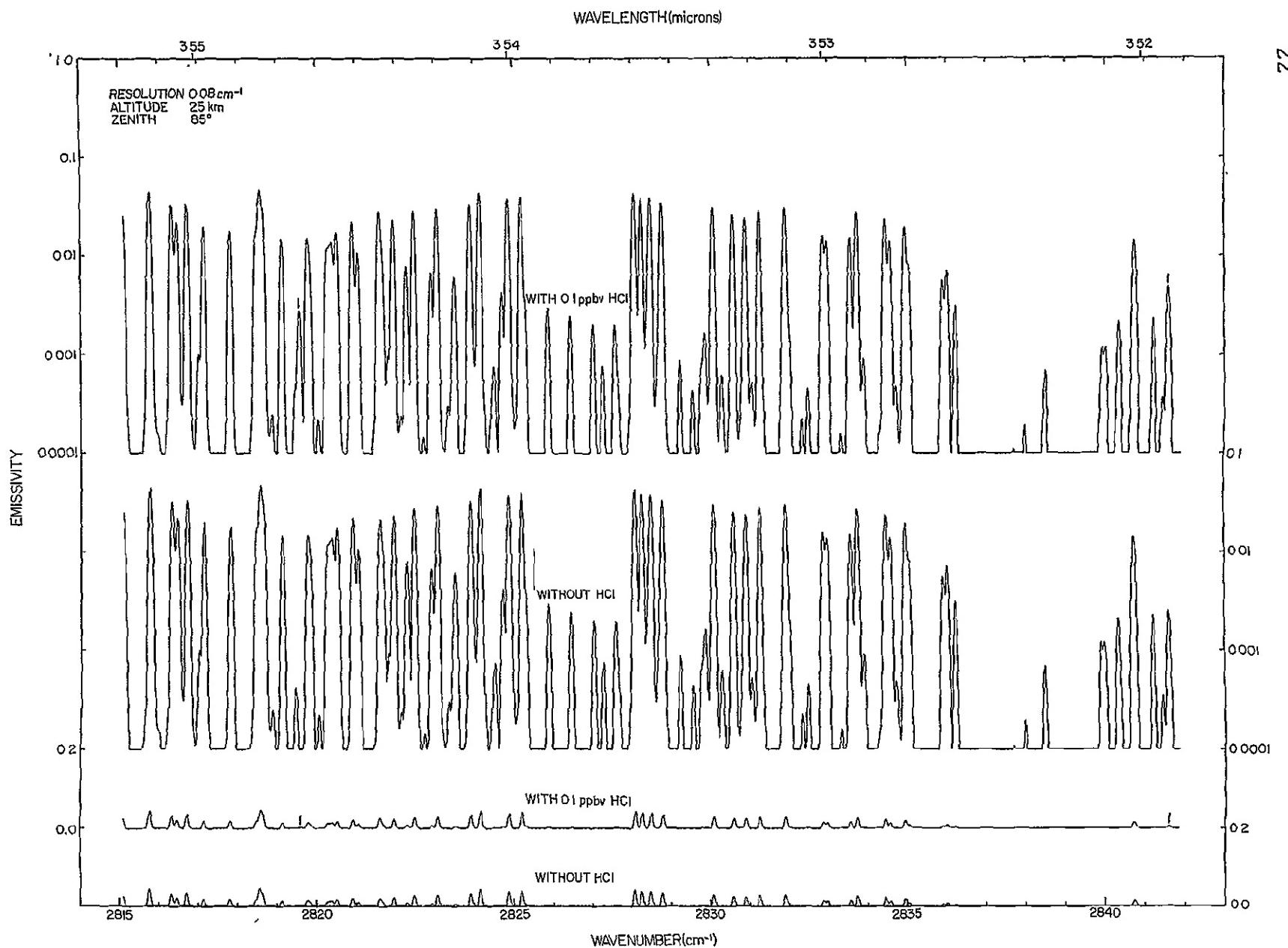


Figure 14.

III. C. HF

The spectral lines in the (0-1) fundamental band of HF are significantly stronger than the (0-1) HCl lines, but in general, the atmospheric interference is stronger in the HF case. The major interferences result from the H_2O and CO_2 $2.7\mu m$ bands and thus the lower wavelength side of these bands is a more promising region for HF detection.

"Infinite" resolution absorption spectra are calculated for a 100-meter path at 5.5 km altitude in Figure 15. Scaling the figures to $\sim 1\%$ emissivity shows that ~ 3 ppbv HF should be measurable in the 4039 cm^{-1} region, where the HF line is significantly stronger than the background atmospheric lines, but is not completely isolated.

The effect of adding 0.2 ppbv HF to the atmosphere on the spectral radiance calculated from 15 km, at 80° zenith angle and 1 cm^{-1} resolution is shown in Figures 16 and 17. Several quite isolated HF features are seen to be comparable in intensity to the atmospheric peaks. The radiance levels of $\sim 2 \times 10^{-12} w\text{ cm}^{-2}\text{ sr}^{-1}\mu m^{-1}$ are marginal for the capability of "state of the art" spectral radiometers.

In terms of atmospheric absorption measurements, Figures 16 and 17 yield $\sim 1\%$ absorptivity. Thus ~ 0.2 ppbv HF should be measurable in the solar spectrum with currently available solar spectrometers which are capable of better than 0.5 cm^{-1} resolution.

Atmospheric absorptions with and without HF are calculated for solar spectrum measurements from an altitude of 25 km, an 85° zenith angle and 0.08 cm^{-1} resolution (Figure 18). The HF line near 4039 cm^{-1} is well above the atmospheric lines and an $\sim 1\%$ measurable absorptivity permits the measurement of ~ 0.02 ppbv HF.

The HF line parameters and the atmospheric line parameters are relatively accurate for this region; however, closer attention should be given to the effects of solar lines in solar spectrum measurements.

High-altitude solar spectra in this region under $\sim 1 \text{ cm}^{-1}$ resolution⁽²⁹⁾ show strong solar features between 4020 and 4030 cm^{-1} and much weaker features near the 4039 cm^{-1} HF line.

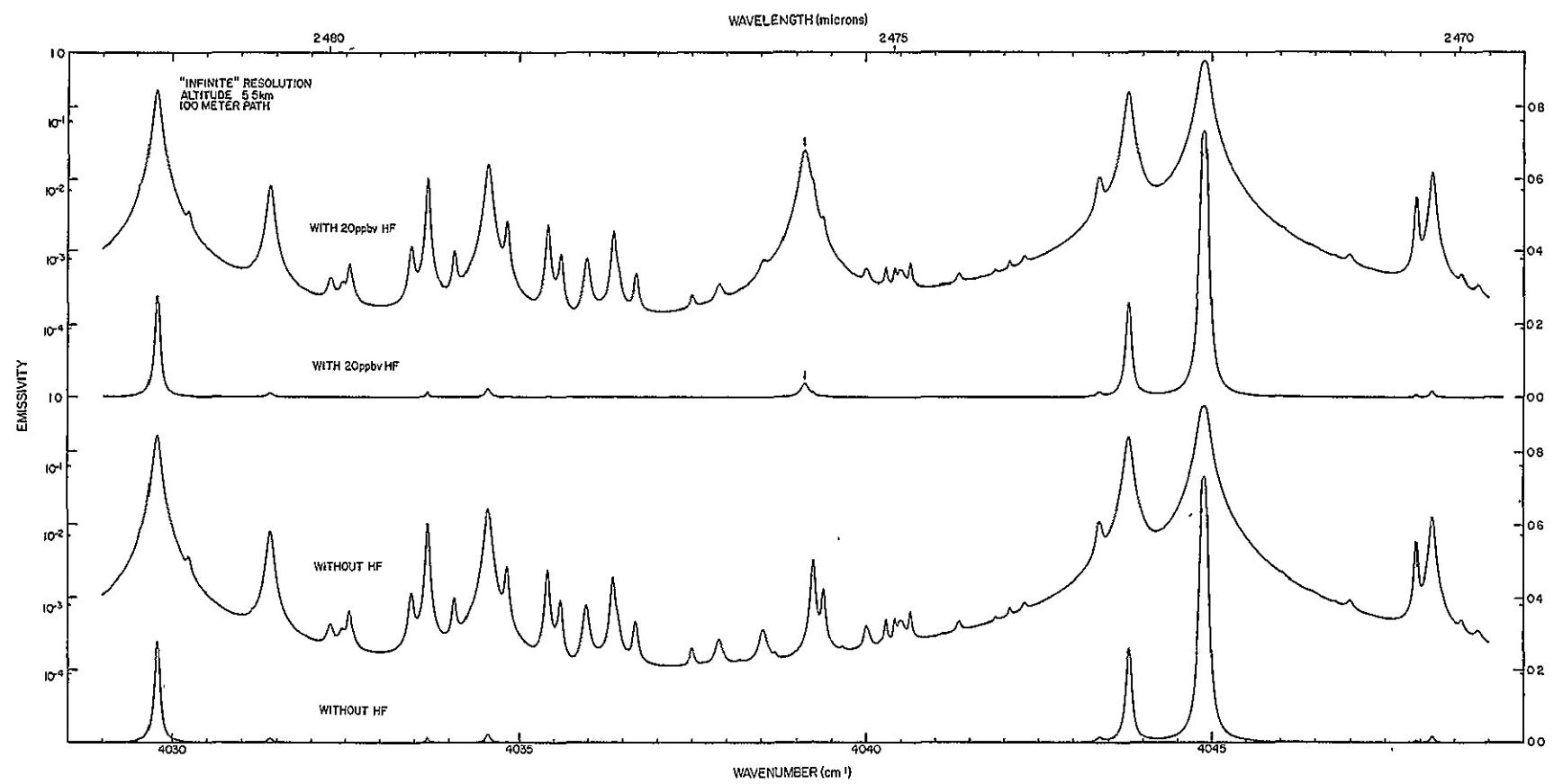


Figure 15.

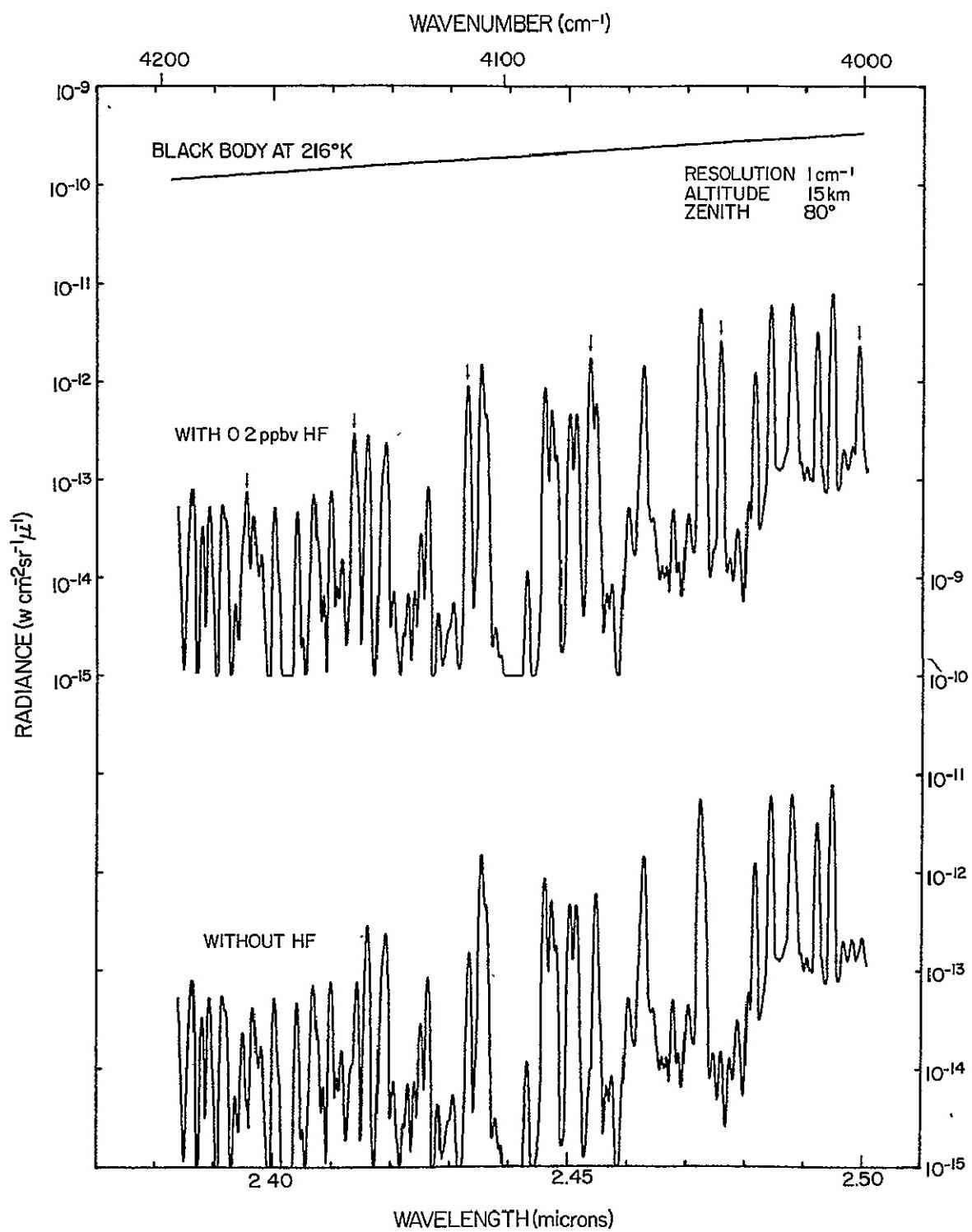


Figure 16.

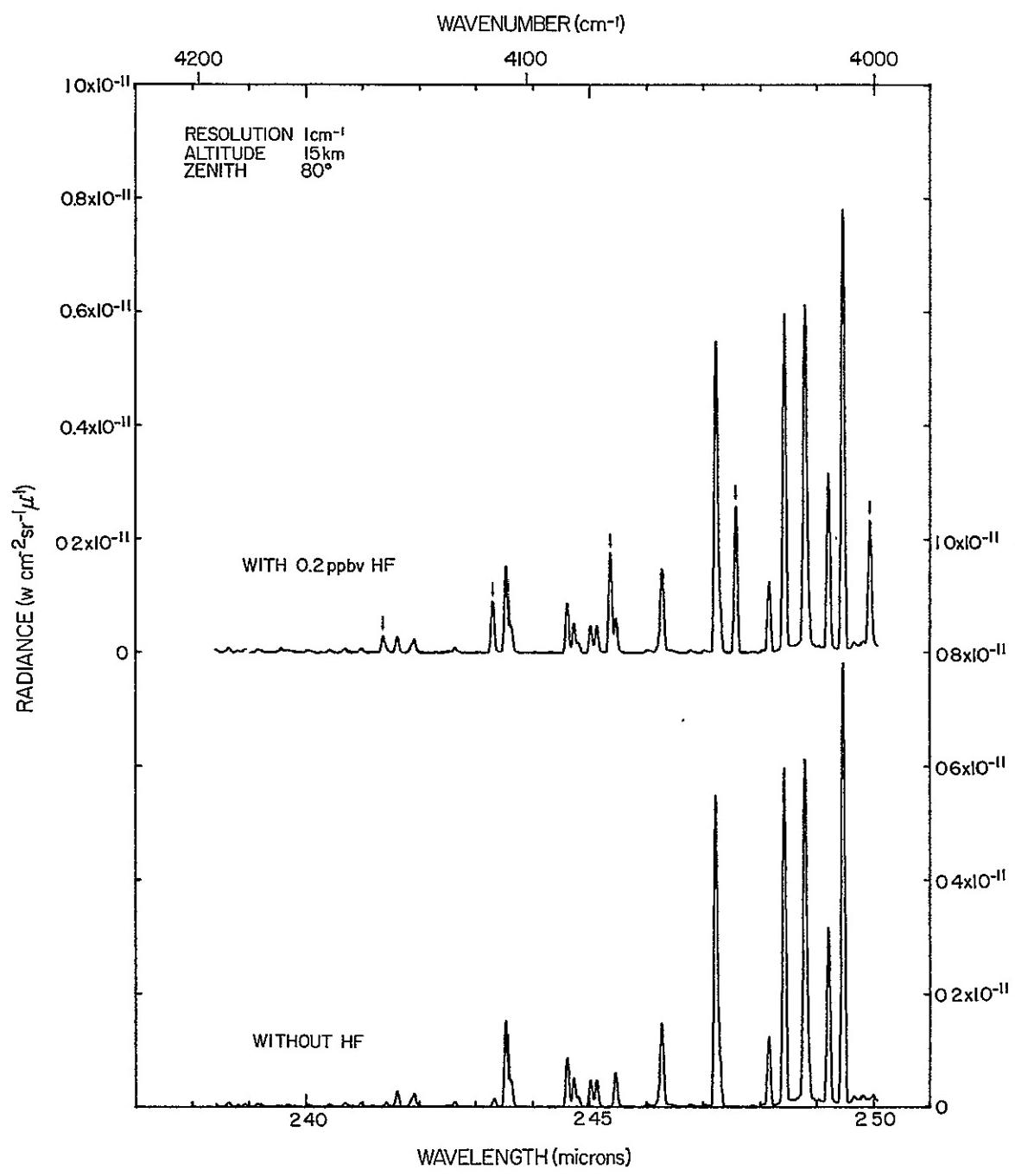


Figure 17.

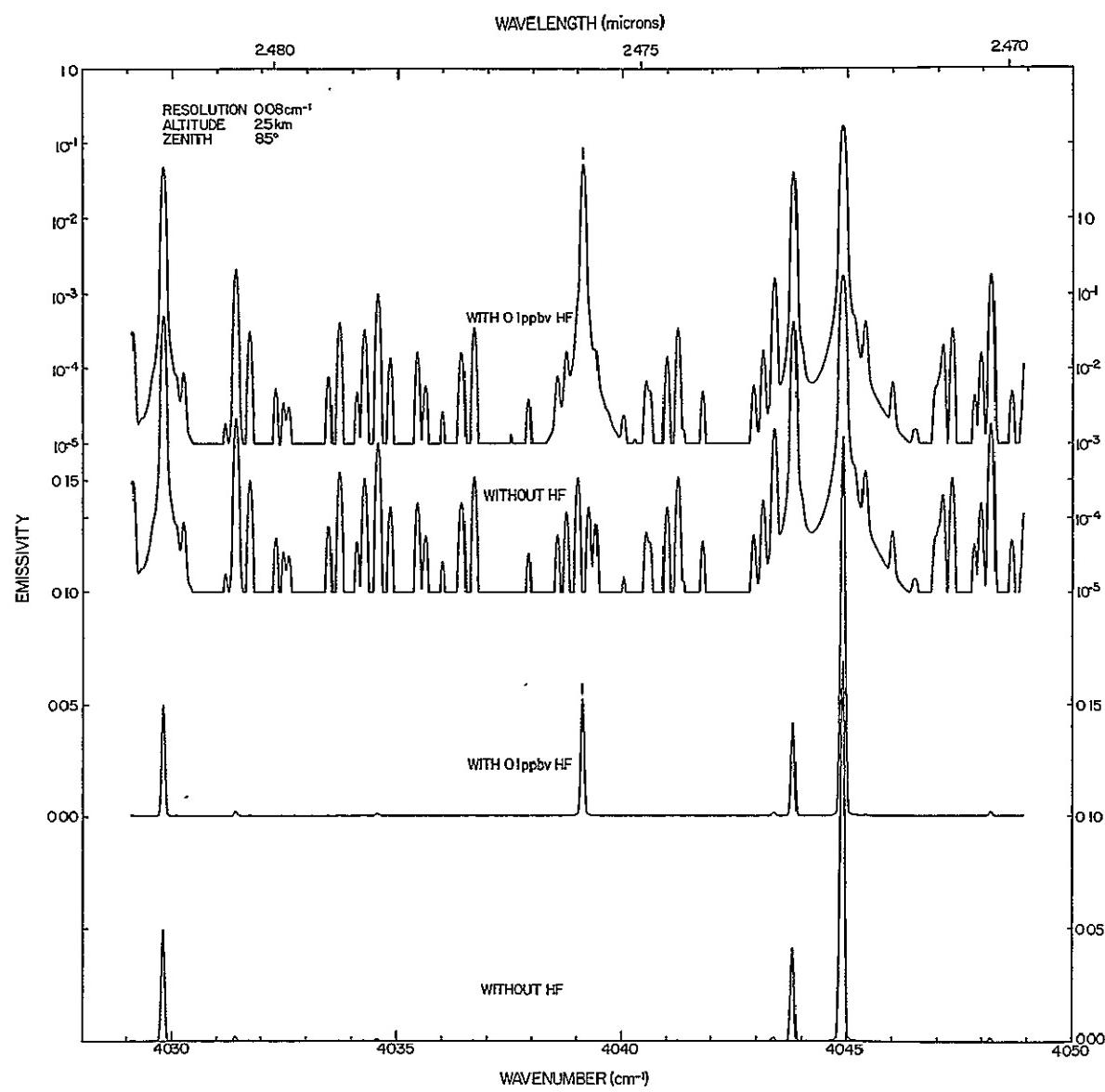


Figure 18.

III. D. HNO_3

A recent paper⁽⁴⁰⁾ has summarized the spectral parameters available for the HNO_3 vibration-rotation bands. Line parameters are not yet available, except for a preliminary compilation for the ν_9 band in the $21.8\mu\text{m}$ region. However the agreement with observed data in this region is not satisfactory yet. Spectral band model parameters are available for the four strong absorption regions at $5.3\mu\text{m}$, $7.5\mu\text{m}$, $11.3\mu\text{m}$ and $21.8\mu\text{m}$, but these are limited to spectral resolutions of 5 to 10 cm^{-1} .

Under these limitations it is more useful to estimate limits of detectability for HNO_3 from the already published stratospheric absorption and emission HNO_3 spectra. From these spectra it is apparent that 3 of the 4 regions listed above are very effective for HNO_3 measurements.

Atmospheric HNO_3 absorption spectra in the $11.3\mu\text{m}$ region at $\sim 0.4 \text{ cm}^{-1}$ resolution are presented in Figure 2 of Ref. 32. From the upper spectra in this figure it can be estimated that at a 90° zenith angle, 30 km altitude, and resolution of 0.08 cm^{-1} , $\sim 1 \text{ ppbv}$ could be measured. The $11.3\mu\text{m}$ region is the most suitable region for HNO_3 measurements above the troposphere, where the interference with the H_2O and CO_2 lines is quite small. Similar estimates for 90° zenith angle and 30 km altitude can also be derived from the $21.8\mu\text{m}$ solar spectra, such as those presented in Figures 2 and 3 of Ref. 36, and from the $7.5\mu\text{m}$ solar spectra presented in Figure 5 of Ref. 32. However the interference with other atmospheric species in the $5.9\mu\text{m}$ and the $21.8\mu\text{m}$ region is greater than in the $11.3\mu\text{m}$ region. The interference in the $5.9\mu\text{m}$ region (see Figure 8 of Ref. 32) makes this region considerably more difficult for HNO_3 measurements. Figure 9 of Ref. 41 shows atmospheric HNO_3 absorption spectra obtained in the $7.5\mu\text{m}$

region at $\sim 0.1 \text{ cm}^{-1}$ resolution, 16 km altitude and zenith angles of 85° to 90° . From these spectra it can be estimated that the $7.5 \mu\text{m}$ region at 0.08 cm^{-1} resolution could allow the measurement of $\sim 5 \text{ ppbv HNO}_3$ from solar spectra taken at an 85° zenith angle and 16 km altitude.

Atmospheric HNO_3 emission spectra in the $11.3 \mu\text{m}$ region, with $\sim 5 \text{ cm}^{-1}$ resolution, such as those shown in Figure 1 of Ref. 42, are available for various altitudes with a zenith angle of 45° . These yield HNO_3 mixing ratios in the range of 1 to 10 ppbv in the 10 to 30 km altitude range. Higher resolution emission spectra will allow a decrease in the detectability limit linearly and also permit HNO_3 measurements at lower altitudes where the interference with H_2O and CO_2 lines is significant. Similar estimates can be made on the basis of the $21.8 \mu\text{m}$ atmospheric spectra such as those presented in Figure 5 of Ref. 36.

III. E. NH₃

The strong ν_2 fundamental band of NH₃ extends over a wide spectral region from ~ 700 to $\sim 1300\text{ cm}^{-1}$. However, many of its strong peaks occur in the strong ν_3 ozone band and will be completely masked in the lower stratosphere between 980 and 1070 cm^{-1} . Therefore two spectral regions, on both sides of the ozone bands, were selected for the present study.

Atmospheric spectral radiance calculated with and without NH₃ at these spectral regions for 15 km, an 80° zenith angle and 1 cm^{-1} resolution, are shown in Figures 19-22. A number of measurable NH₃ features can be seen in the 10.3 to $12.4\mu\text{m}$ region (Figures 19 and 20) with 0.1 ppbv NH₃. The radiance levels are above $10^{-7}\text{ w cm}^{-2}\text{ sr}^{-1}\mu\text{m}^{-1}$, well within the capability of current spectral radiometers. Figures 21 and 22 cover the 7.9 to $9.3\mu\text{m}$ region where the many interfering background lines require larger minimum detectable amounts of NH₃. Also, uncertainties are apparent with regard to the $8.6\mu\text{m}$ N₂O band in this region.

It should be noted that while the atmospheric H₂O and CO₂ line parameters, as well as the NH₃ line parameters in the 9 - $12\mu\text{m}$ region, are relatively accurate, the $11.3\mu\text{m}$ HNO₃ band was not included in the present calculation. However, the HNO₃ features are much broader and thus would allow observation of the NH₃ peaks superimposed on the HNO₃ spectra. The interference of NH₃ lines with both the CO₂ and HNO₃ lines are minimal near 930 cm^{-1} , (see also Ref. 32) which is thus a promising region for NH₃ detection.

The interference effects are further reduced on absorption spectra calculated for solar spectrum measurements from an altitude

of 25 km, an 85° zenith angle and 0.08 cm^{-1} resolution (Figure 23). Here a number of isolated NH_3 peaks make it feasible to measure $\sim 0.05 \text{ ppbv } \text{NH}_3$. The interference effects are similar for solar spectrum measurements under similar resolution and zenith angles from an altitude of 35 km, as is shown in Figure 24. However, the optical path is reduced in this case and thus requires a larger proportion of NH_3 for its spectral features to become evident.

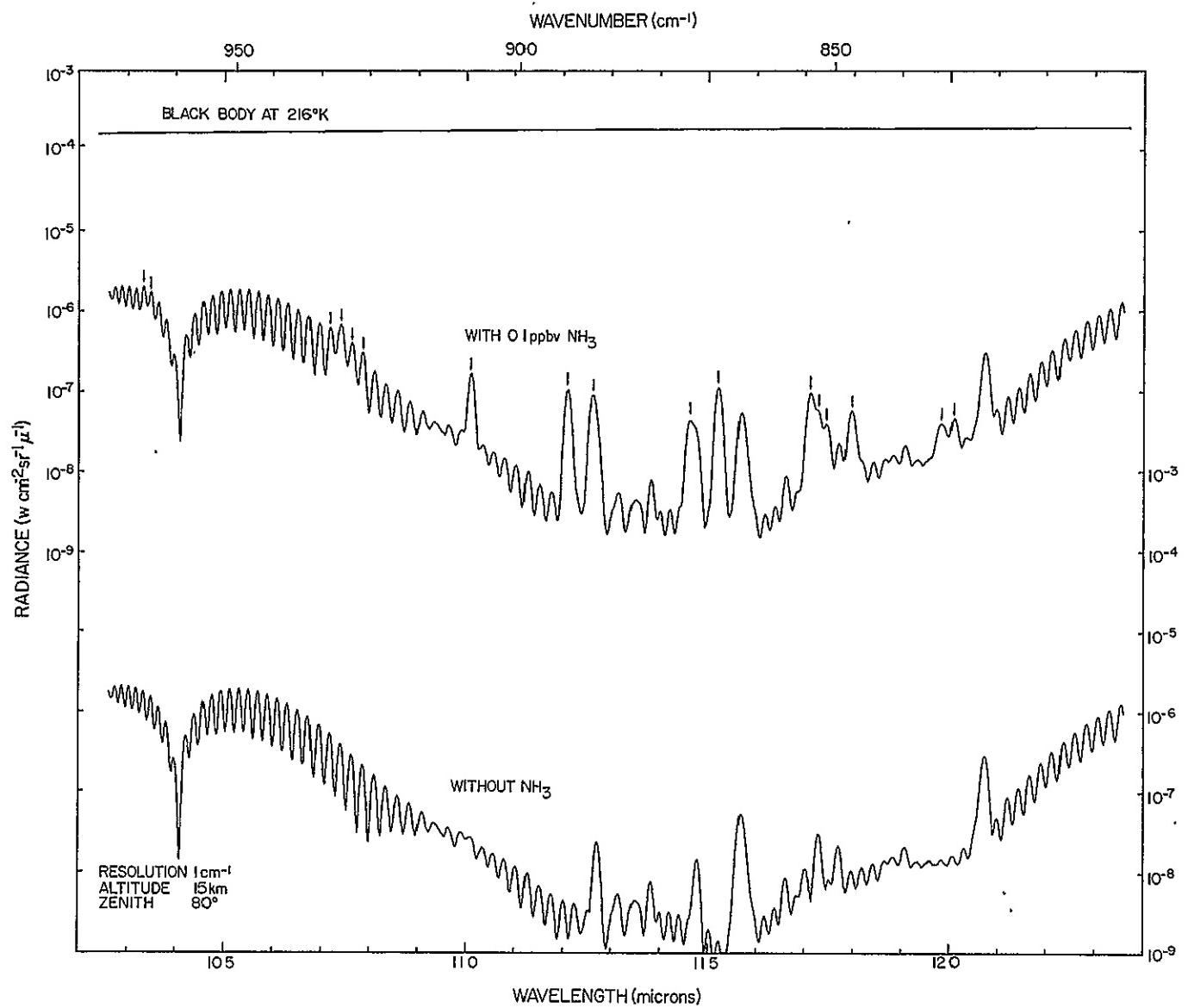


Figure 19.

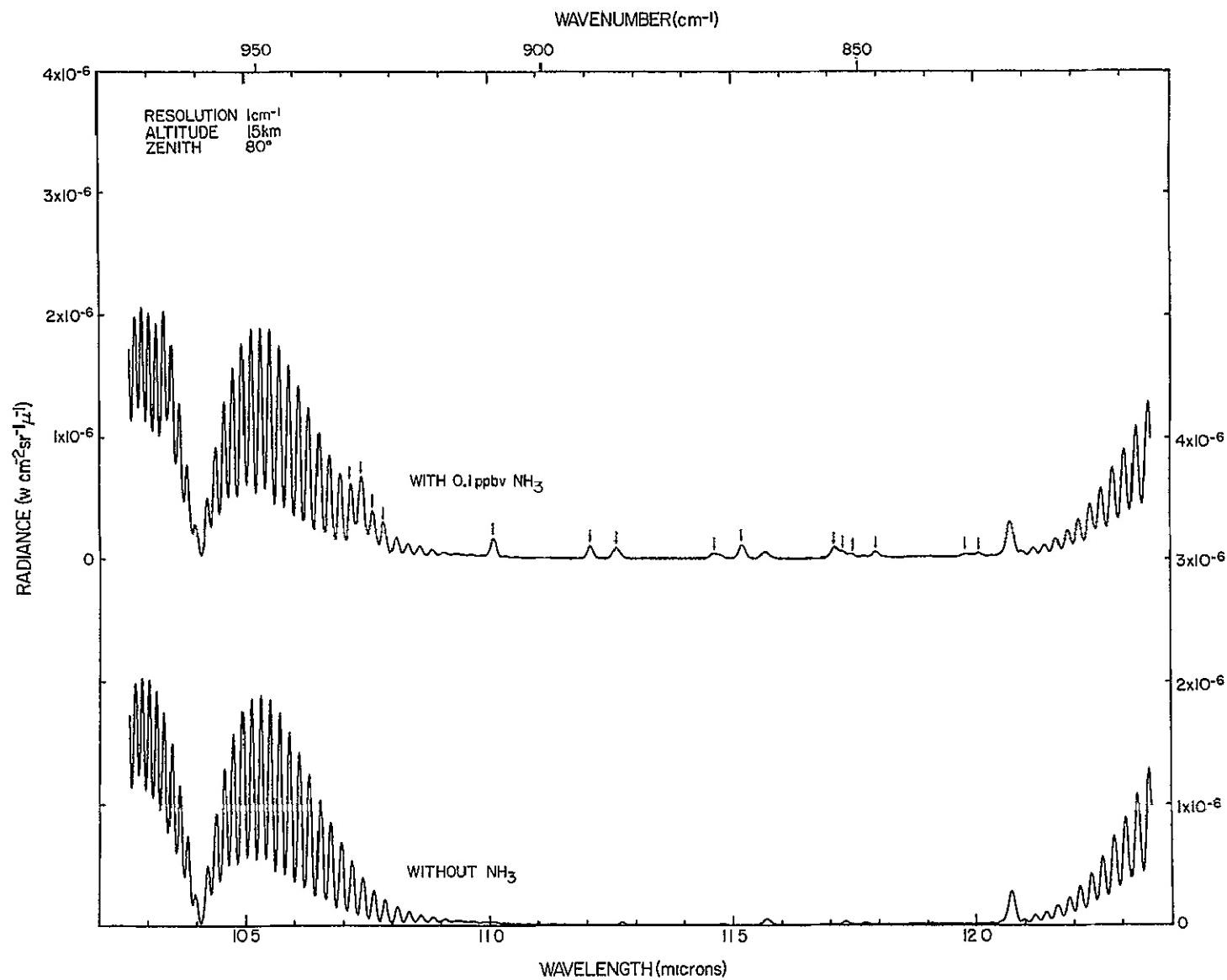


Figure 20.

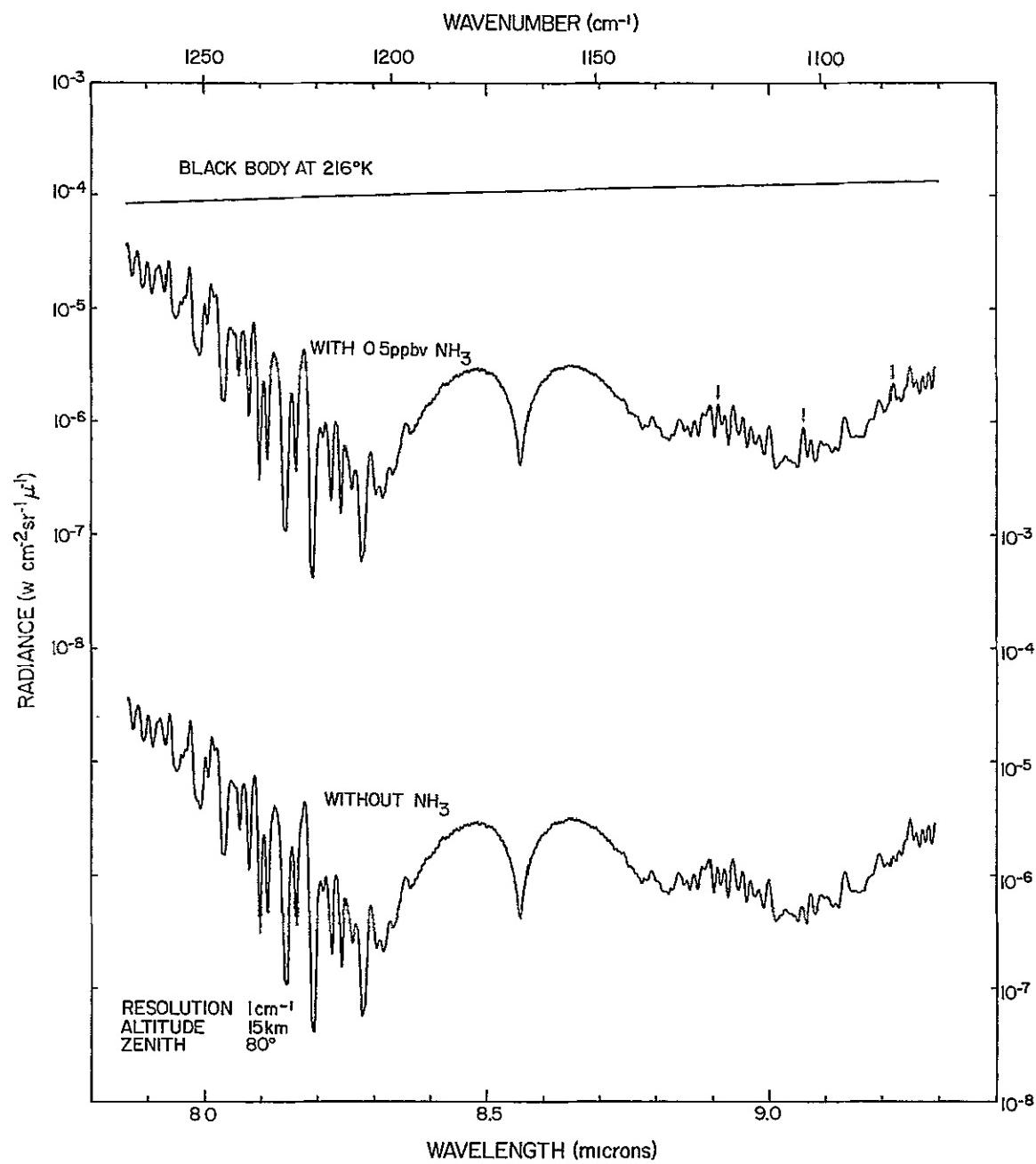


Figure 21.

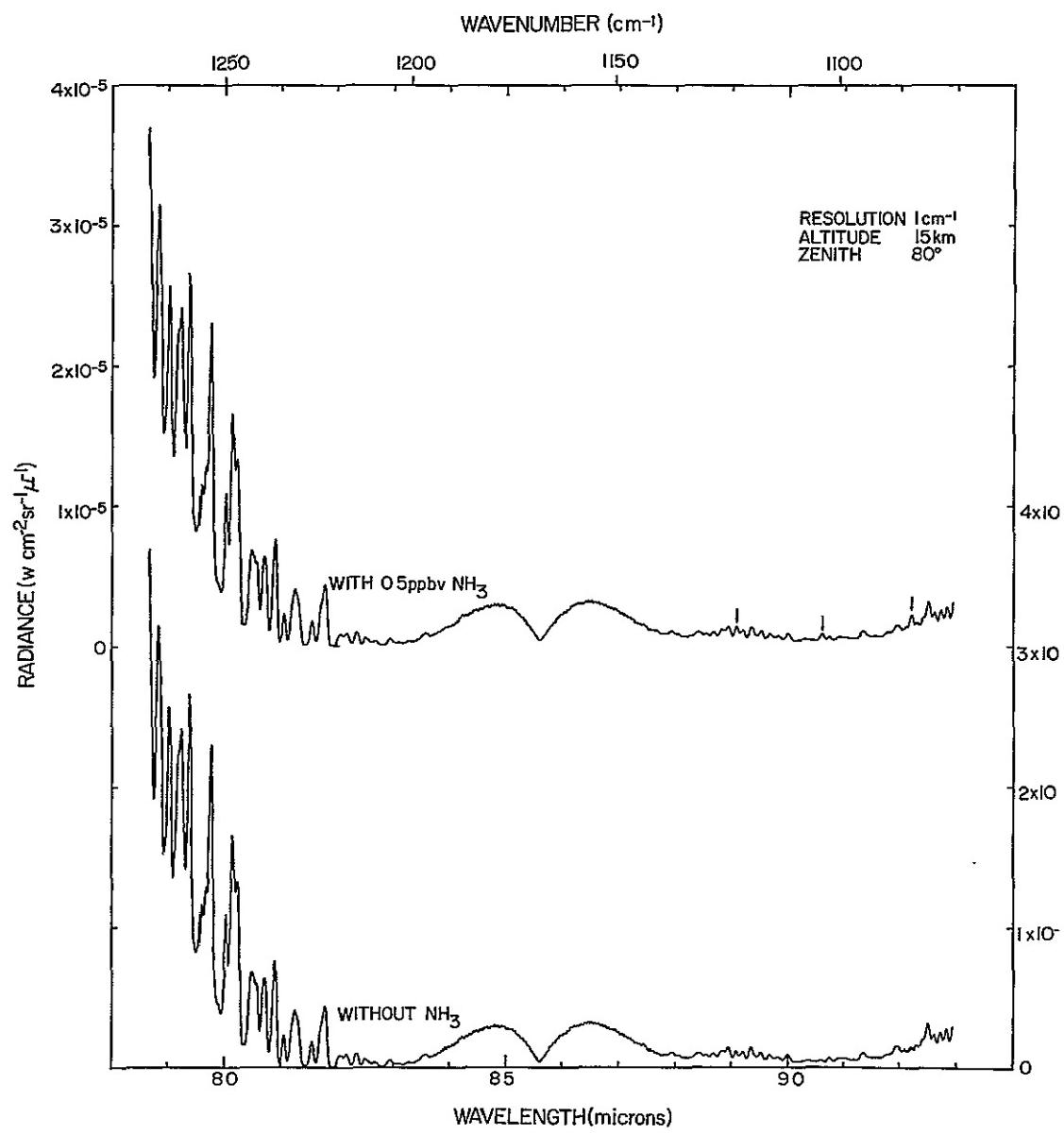


Figure 22.

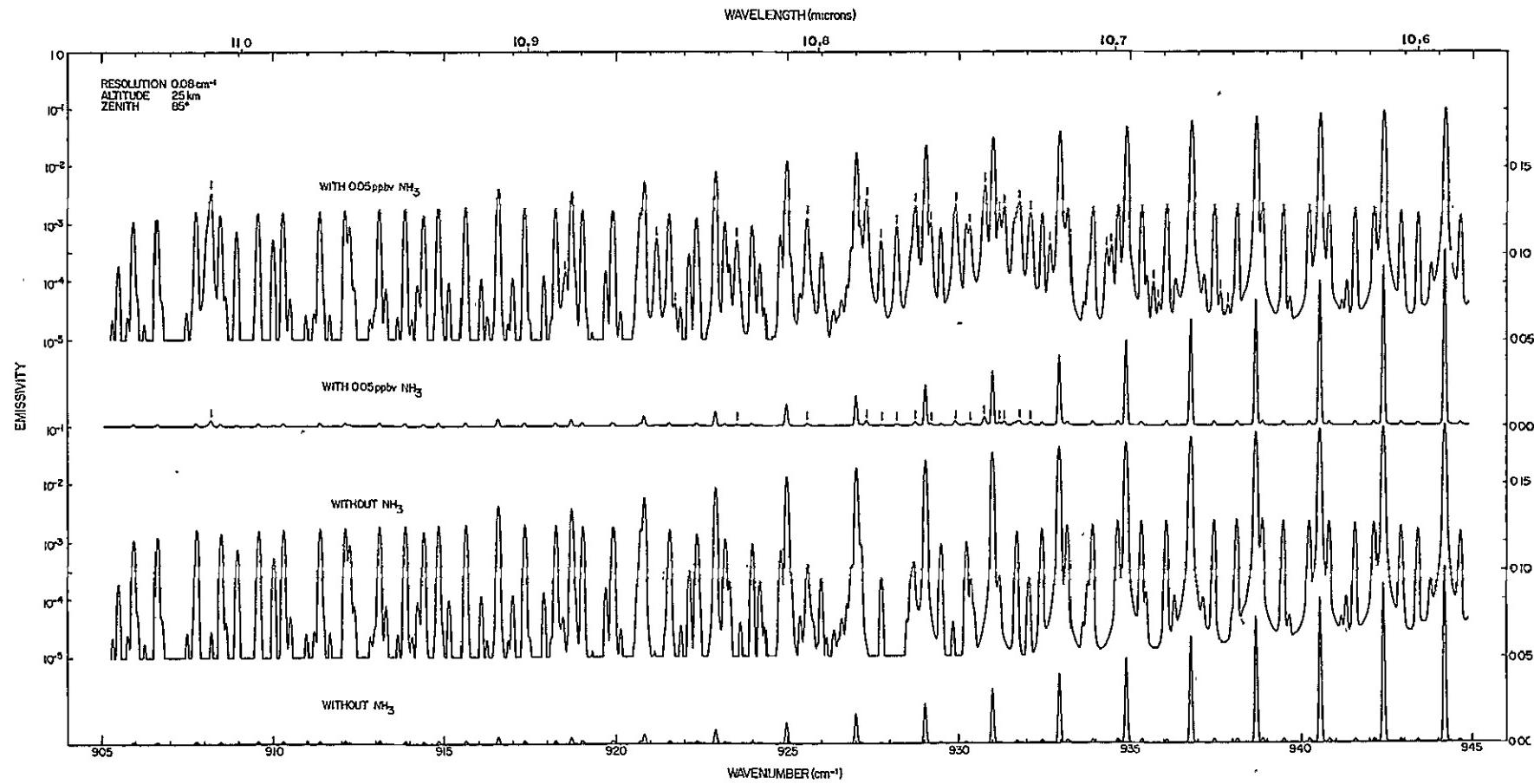


Figure 23.

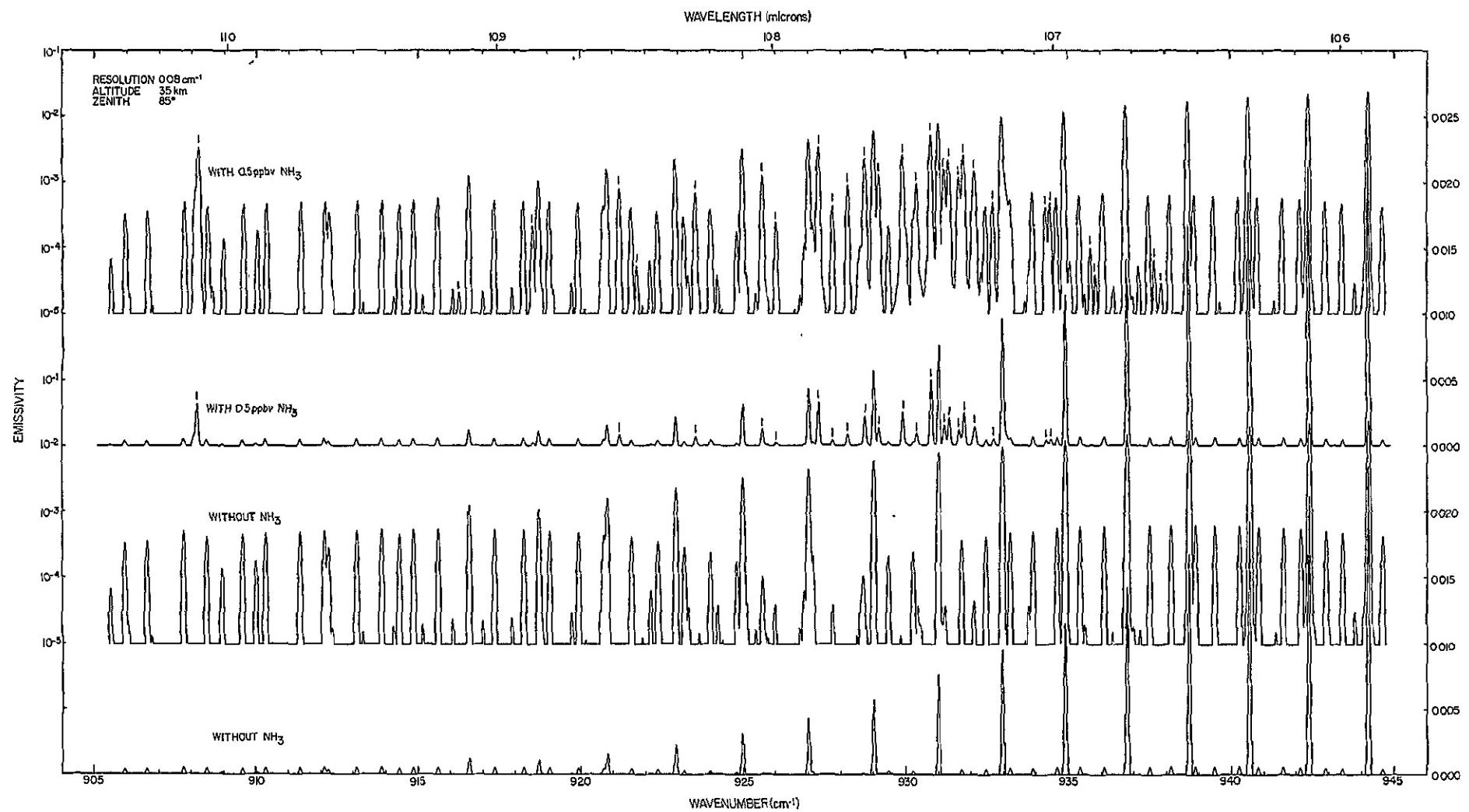


Figure 24.

III. F. NO

A detailed study of possible NO absorption measurements from the solar spectrum as observed from various altitudes has been presented previously.⁽¹⁹⁾ Additional emission and absorption spectra are calculated here.

Atmospheric emission spectra are calculated for a spectral radiometer at an altitude of 15 km, an 80° zenith angle and 0.8 cm⁻¹ resolution in Figures 25 and 26. It is seen that the NO lines overlap with numerous atmospheric lines. The addition of 5 ppbv NO results in several marginally measurable NO features, but none of the NO features is completely isolated from atmospheric lines. The radiance levels are of the order of 10⁻⁷ w cm⁻² sr⁻¹ μm⁻¹ and are within the capability of "state of the art" spectral radiometers. The absorptivities of the NO features shown in Figures 25 and 26 are of the order of ~1%, but do not include the solar CO features.⁽¹⁹⁾

"Infinite" resolution absorption spectra are calculated in Figure 27 for a 100 meter path at 15 km altitude. The atmospheric interference is significantly reduced under these conditions and many of the NO features are quite isolated. However, the optical path is considerably smaller so that the ~1% absorptivity criterion permits the measurement of ~50 ppbv NO (by scaling Figure 27).

It should be noted that while the NO line parameters are relatively accurate, the accuracy of many of the atmospheric line parameters in the NO region is significantly less, especially for the weak CO₂ lines.⁽¹⁹⁾ In addition, the solar CO effects, as presented previously⁽¹⁹⁾ should be taken into account.

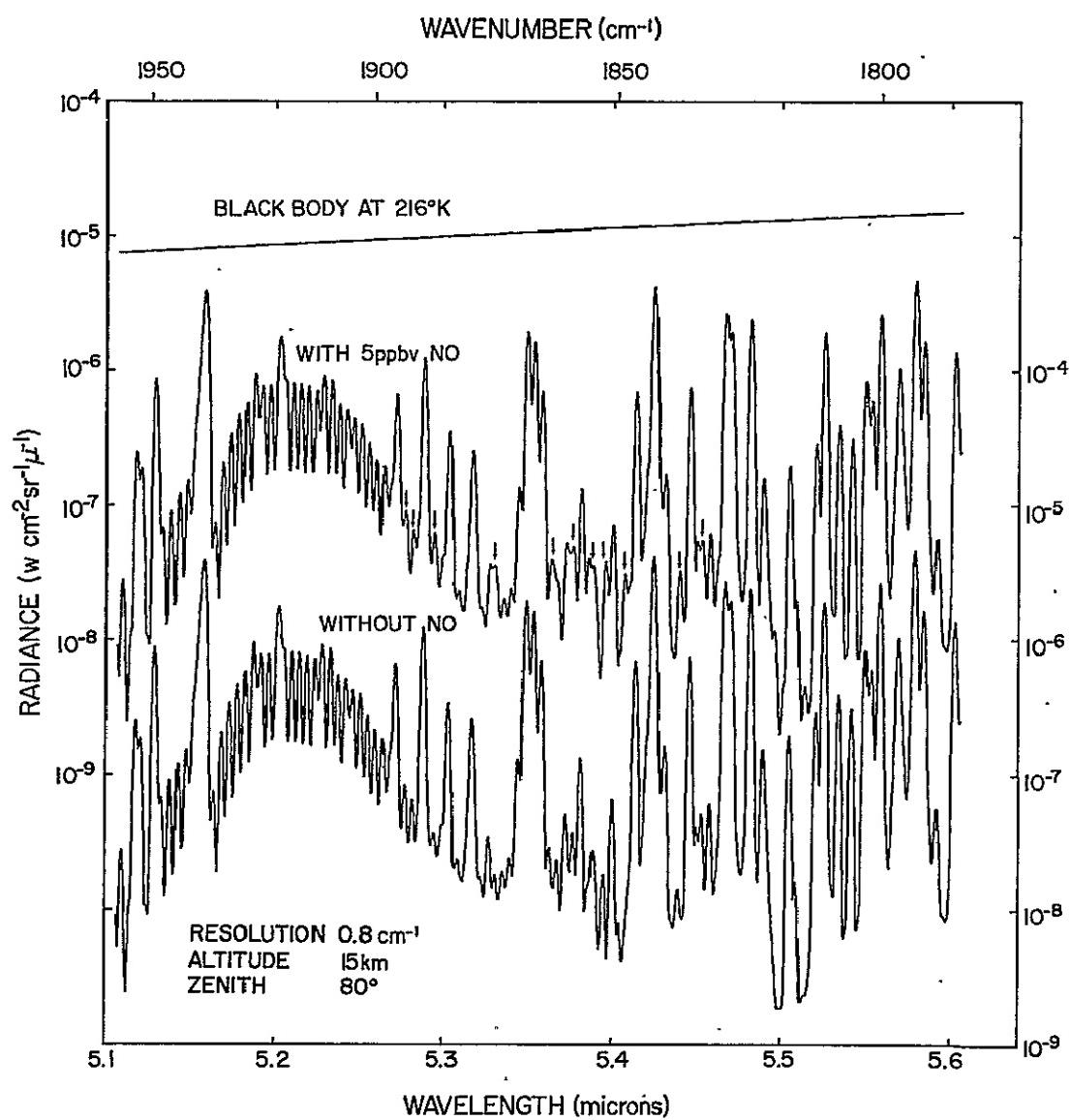


Figure 25.

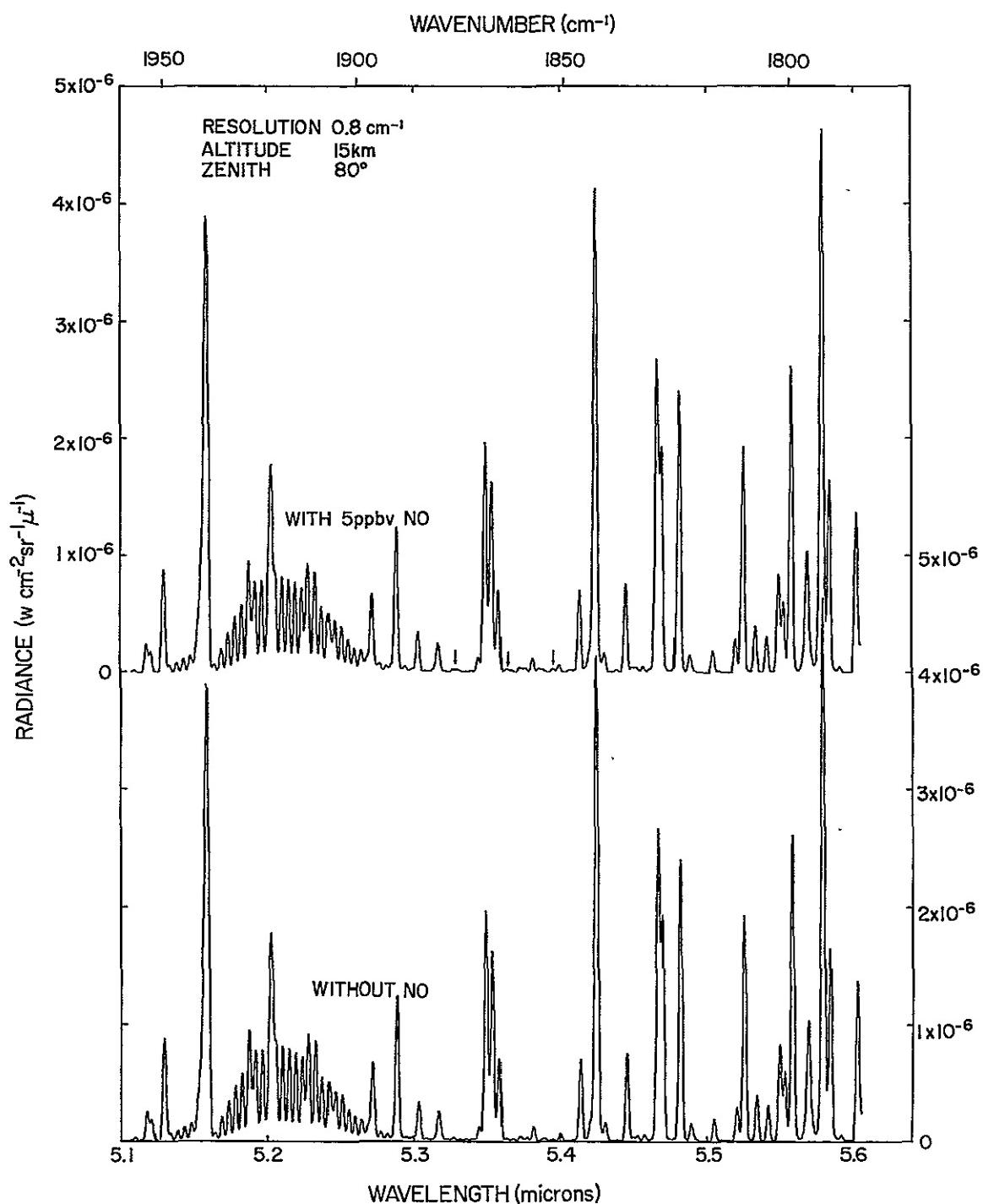


Figure 26.

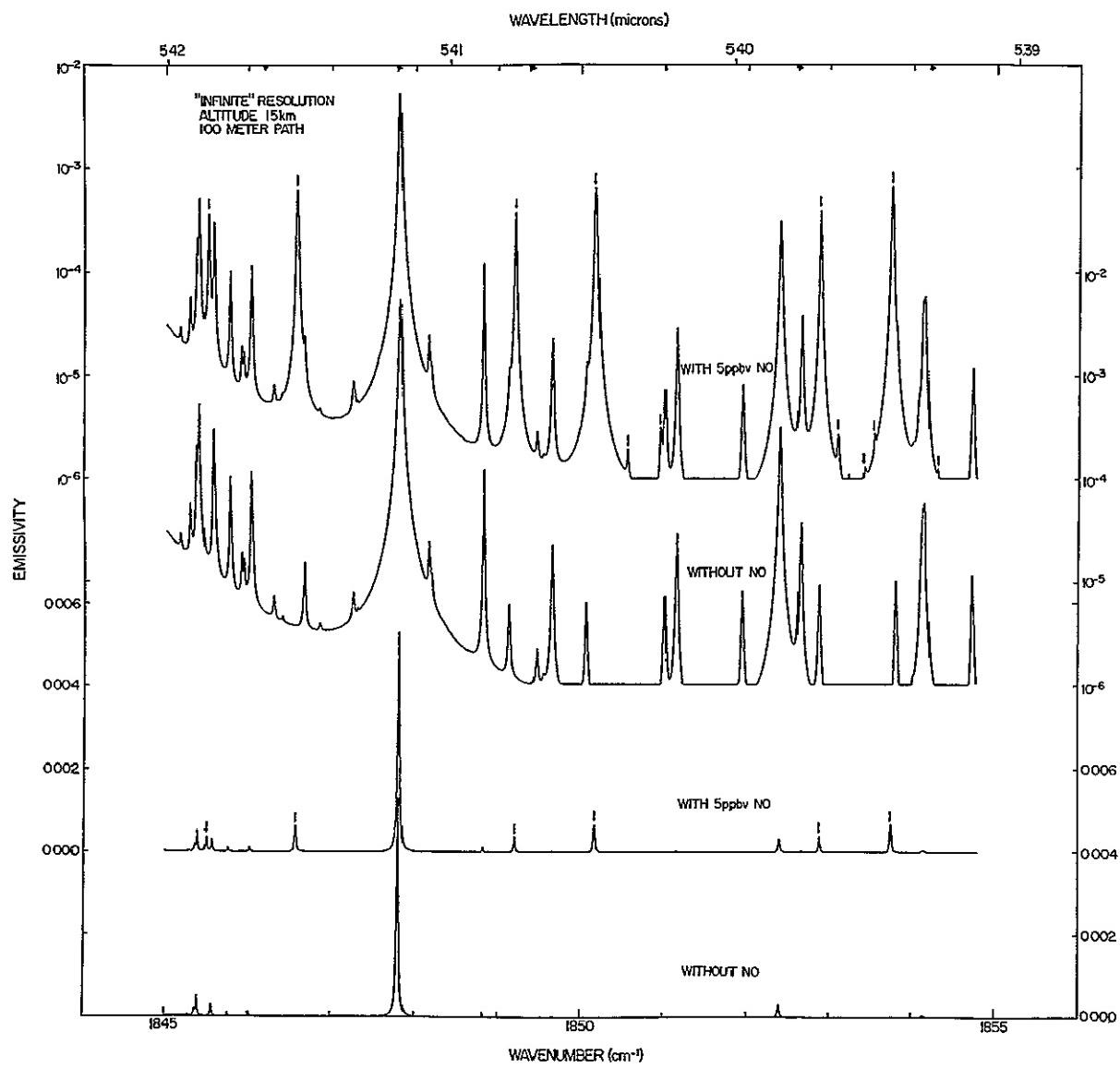


Figure 27.

III. G. NO₂

The ν_3 band of NO₂ is the strongest of the NO₂ bands. However this band occurs near the center of the strong 6.3 μm H₂O band so that the determination of NO₂ from low resolution, long path measurements at low altitudes would be impractical. The masking effect of the numerous atmospheric lines on the NO₂ lines is significantly reduced when higher spectral resolution and higher altitudes are used, thereby making measurements feasible.

Atmospheric absorption calculations, with and without NO₂, are shown in Figures 28 and 29 for a solar spectrum measurement from 25 km altitude, an 85° zenith angle, and 0.08 cm⁻¹ resolution. It is seen that ~1% absorptivity permits the measurement of ~0.1 ppbv NO₂ under these conditions. Many NO₂ lines are clearly isolated from the atmospheric lines, especially near 1605 cm⁻¹ where a strong NO₂ line group occurs in an atmospheric "window." Several isolated NO₂ line groups also occur between H₂O lines near 1612 cm⁻¹.

Similar calculations for 35 km are shown in Figure 30. It is apparent that ~1 ppbv NO₂ should be measurable from many isolated lines of NO₂ with an absorptivity of ~1%.

The H₂O line parameters are known to relatively high accuracy in the NO₂ region. However, uncertainties exist with regard to the CH₄⁽²⁶⁾ lines in this region and also the NO₂ lines used here.

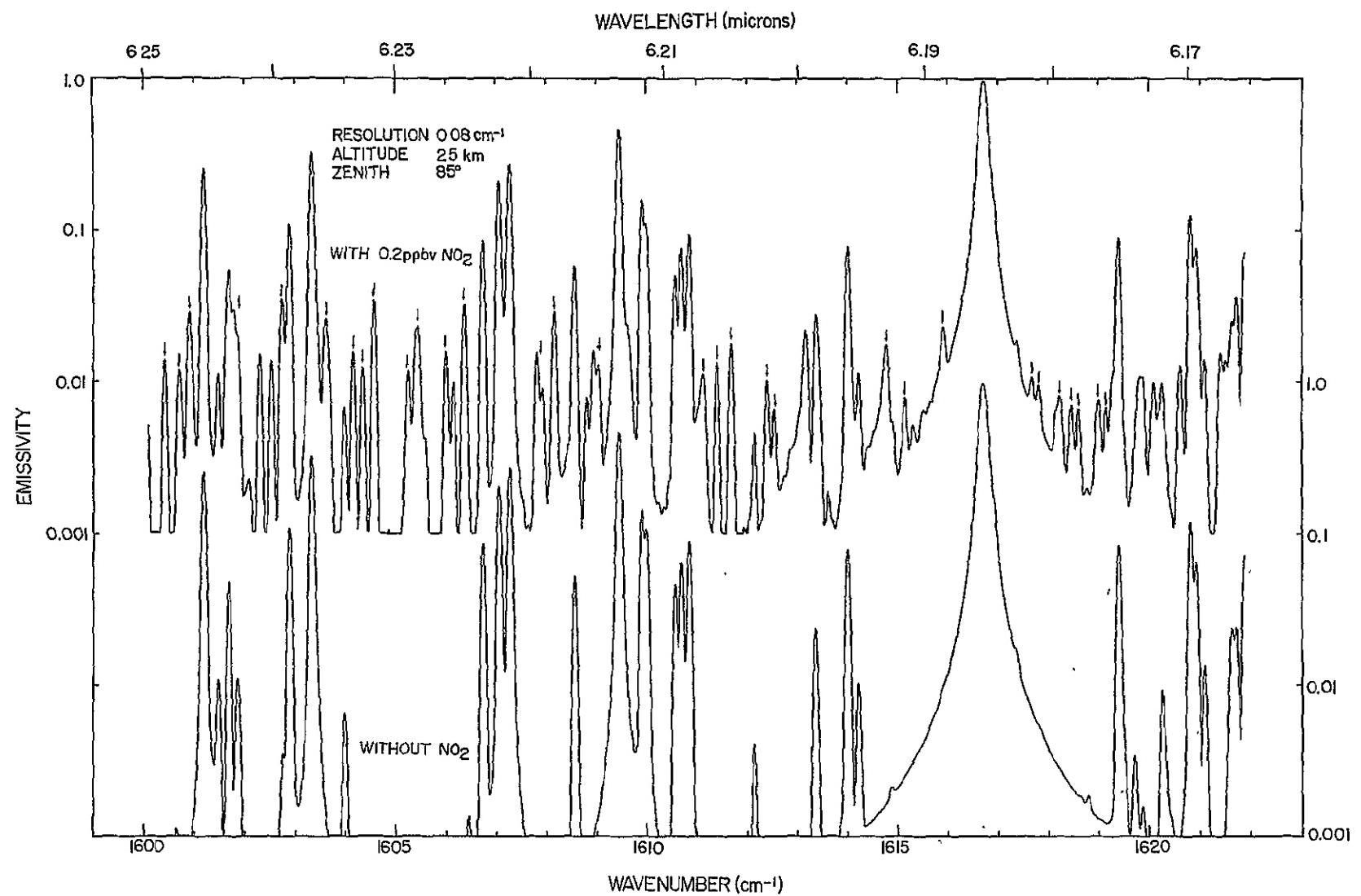


Figure 28.

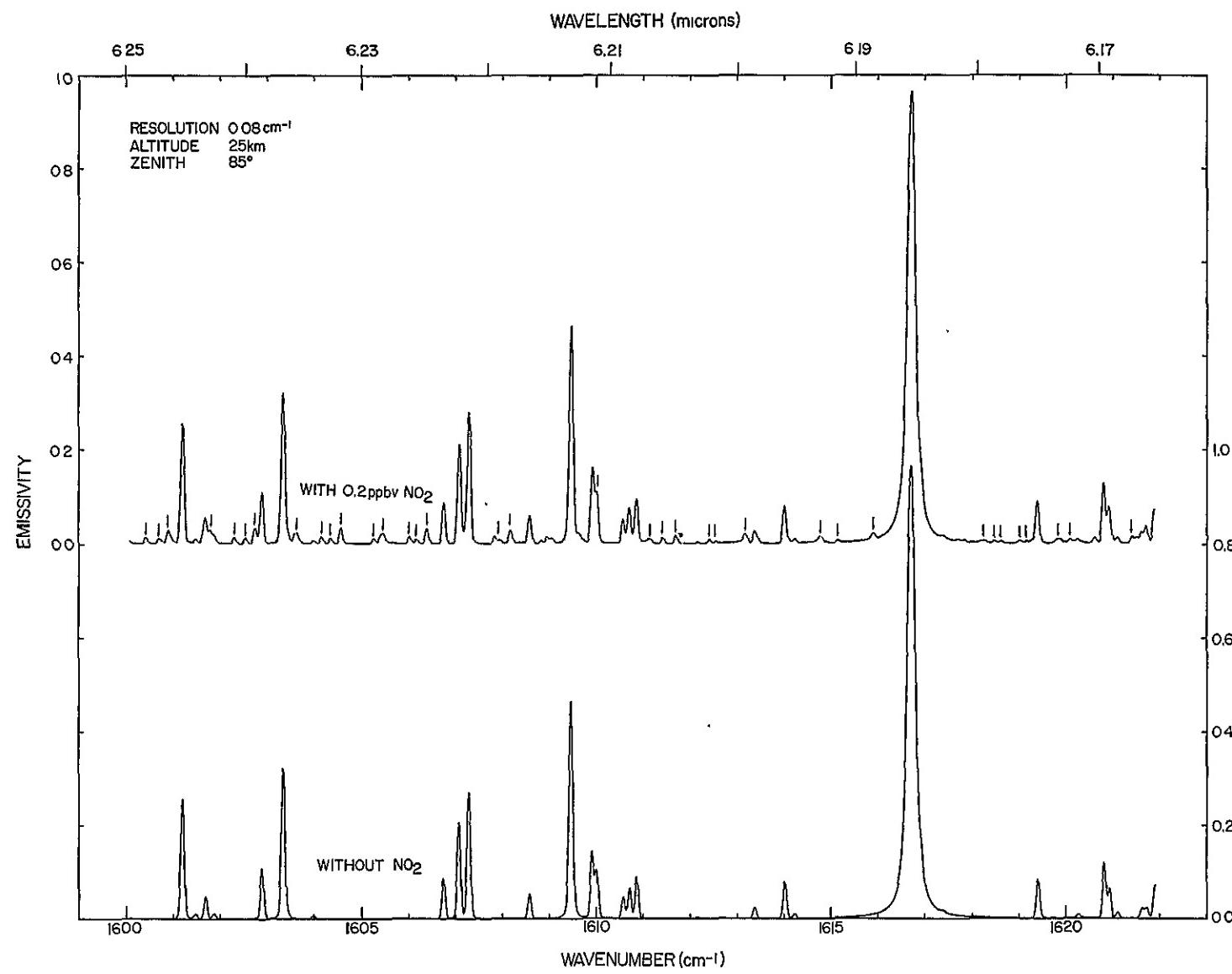


Figure 29.

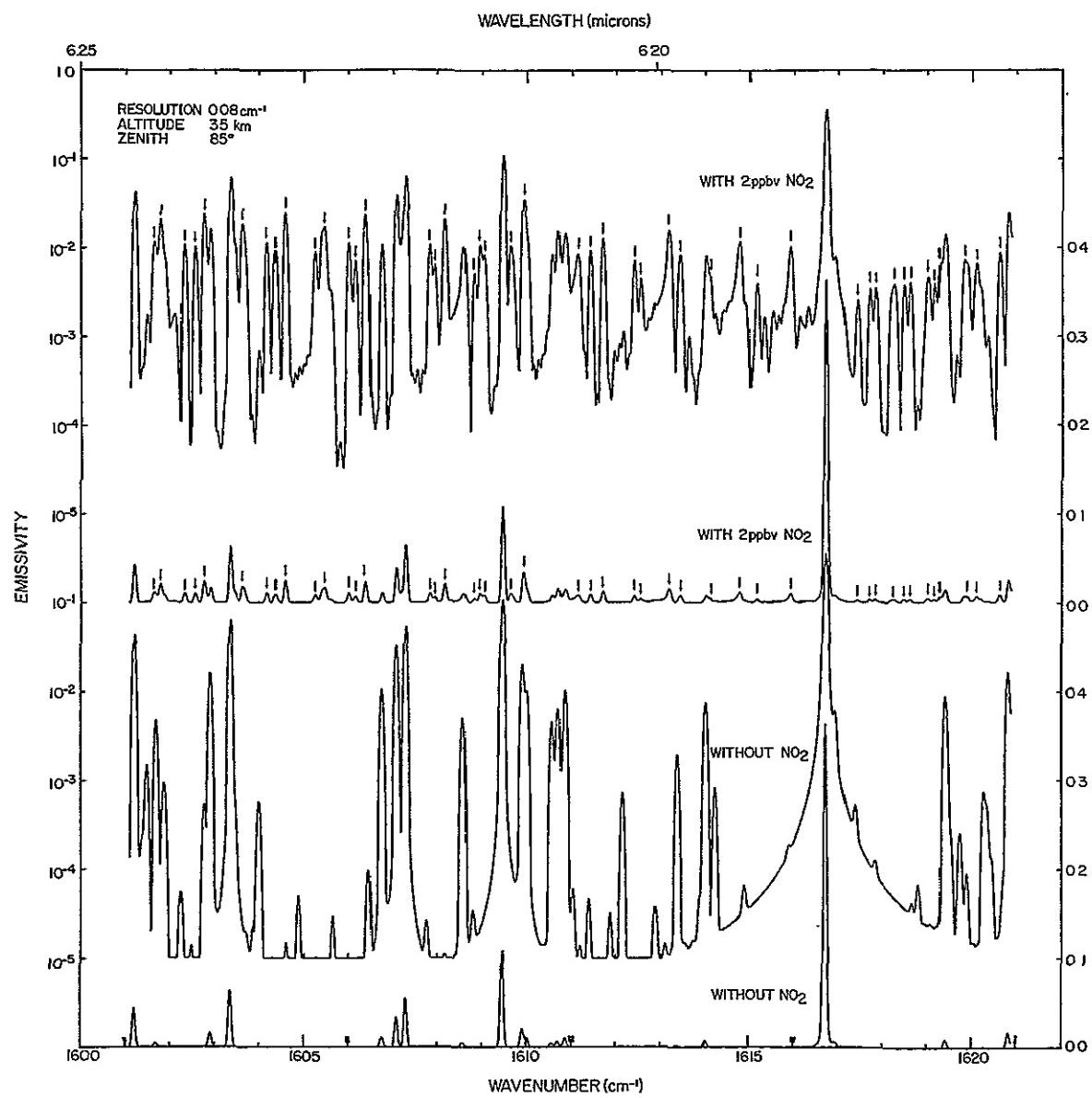


Figure 30.

III. H. SO₂

The ν_3 vibration-rotation fundamental band, with the strong Q-branch near 1360 cm⁻¹, is the strongest of the SO₂ bands.

The atmospheric spectral radiance calculated for a spectral radiometer at 5.5 km altitude, an 80° zenith angle and 1 cm⁻¹ resolution (Figure 31) shows significant signal change with the addition of 10 ppbv SO₂. However, the spectrum is dominated by the background atmosphere and is also close to a blackbody spectrum. As a result no individual features of SO₂ can be isolated and the SO₂ contribution is manifested in small changes of the broad atmospheric peaks. Derivation of SO₂ amounts from atmospheric radiance measurements under such conditions is bound to have a very large uncertainty, despite the calculated large change in signal.

The situation is somewhat improved when similar calculations are made for 15 km altitude in Figure 32, however, most of the SO₂ features are still masked by the background atmospheric lines. The addition of 5 ppbv SO₂ to the atmosphere results in potentially measurable changes of the intensities of the atmospheric peaks.

In both of the above cases the atmospheric radiance levels are larger than 10⁻⁵ w cm⁻² sr⁻¹ μm⁻¹ and the absorptivities are between 10 and 50%. These are well within the capability of current instruments.

The atmospheric absorption spectra in a solar spectrum measurement from an altitude of 25 km, an 85° zenith angle and 0.08 cm⁻¹ resolution show that the 1360 cm⁻¹ region allows the measurement of ~0.5 ppbv SO₂ from ~1% absorptivity of isolated SO₂ peaks (Figures 33 and 34). These SO₂ peaks are a portion of the strong Q-branch of the ν_3 band. The SO₂ peaks near 1365 cm⁻¹ also have potential for SO₂ measurements in the atmosphere.

The spectral region used here occurs at the wings of the 1326 cm^{-1} HNO_3 band (not included in the present calculations) and only a little interference is expected from this band. It should be noted, however, that the line parameters of the $\nu_3 \text{SO}_2$ band have not been studied enough to guarantee the accuracy achieved for most of the other line parameters used in this study.

WAVENUMBER (cm^{-1})

49

1380

1360

1340

RADIANCE ($\text{w cm}^{-2} \text{sr}^{-1} \mu\text{m}^{-1}$)

3×10^{-4}

2×10^{-4}

1×10^{-4}

3×10^{-4}

2×10^{-4}

1×10^{-4}

7.2

7.3

7.4

7.5

WAVELENGTH (microns)

RESOLUTION 1 cm^{-1}
ALTITUDE 5.5 km
ZENITH 80°

BLACK BODY AT 253°K

WITH 10 ppbv SO_2

BLACK BODY AT 253°K

WITHOUT SO_2

Figure 31.

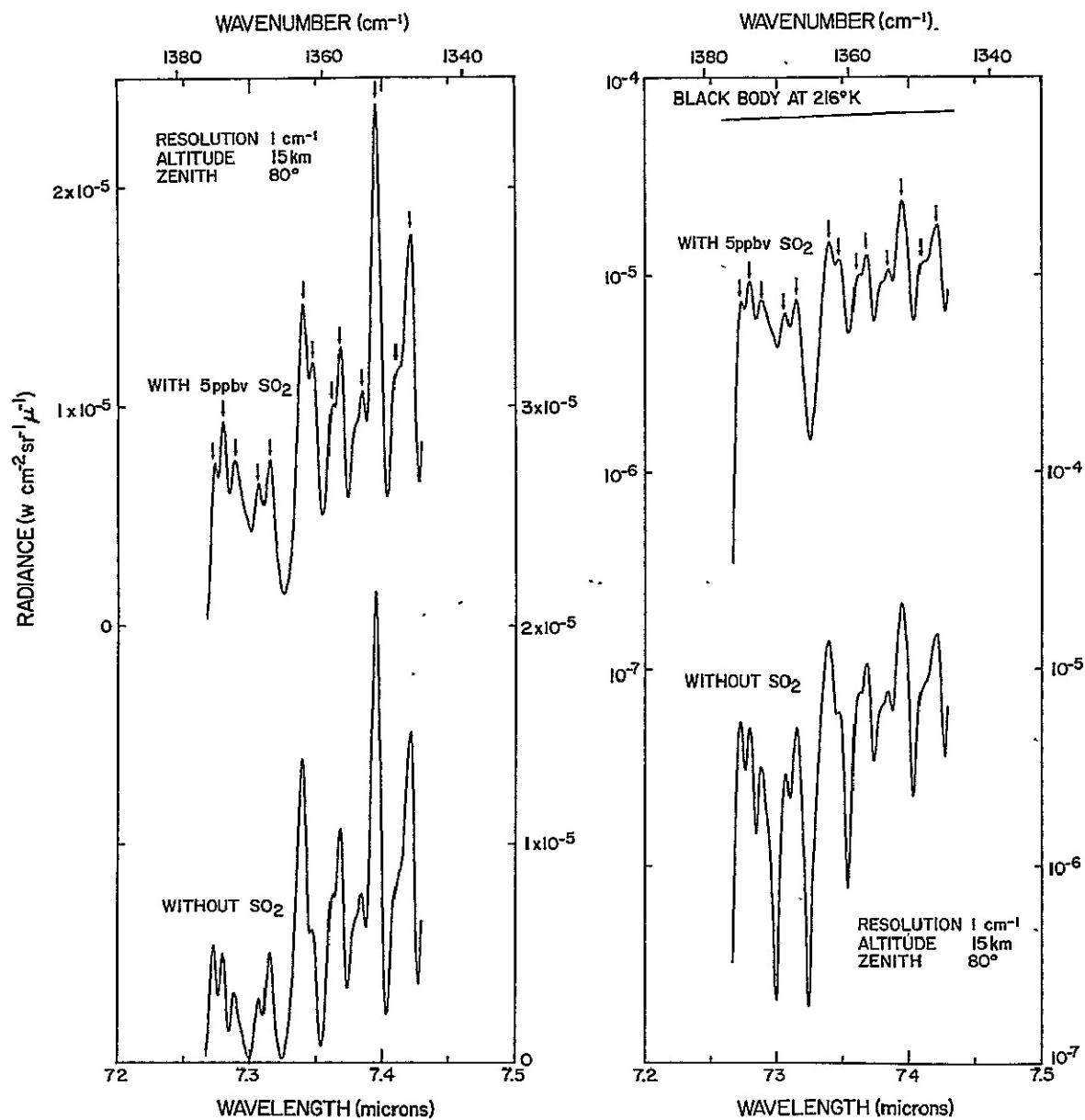


Figure 32.

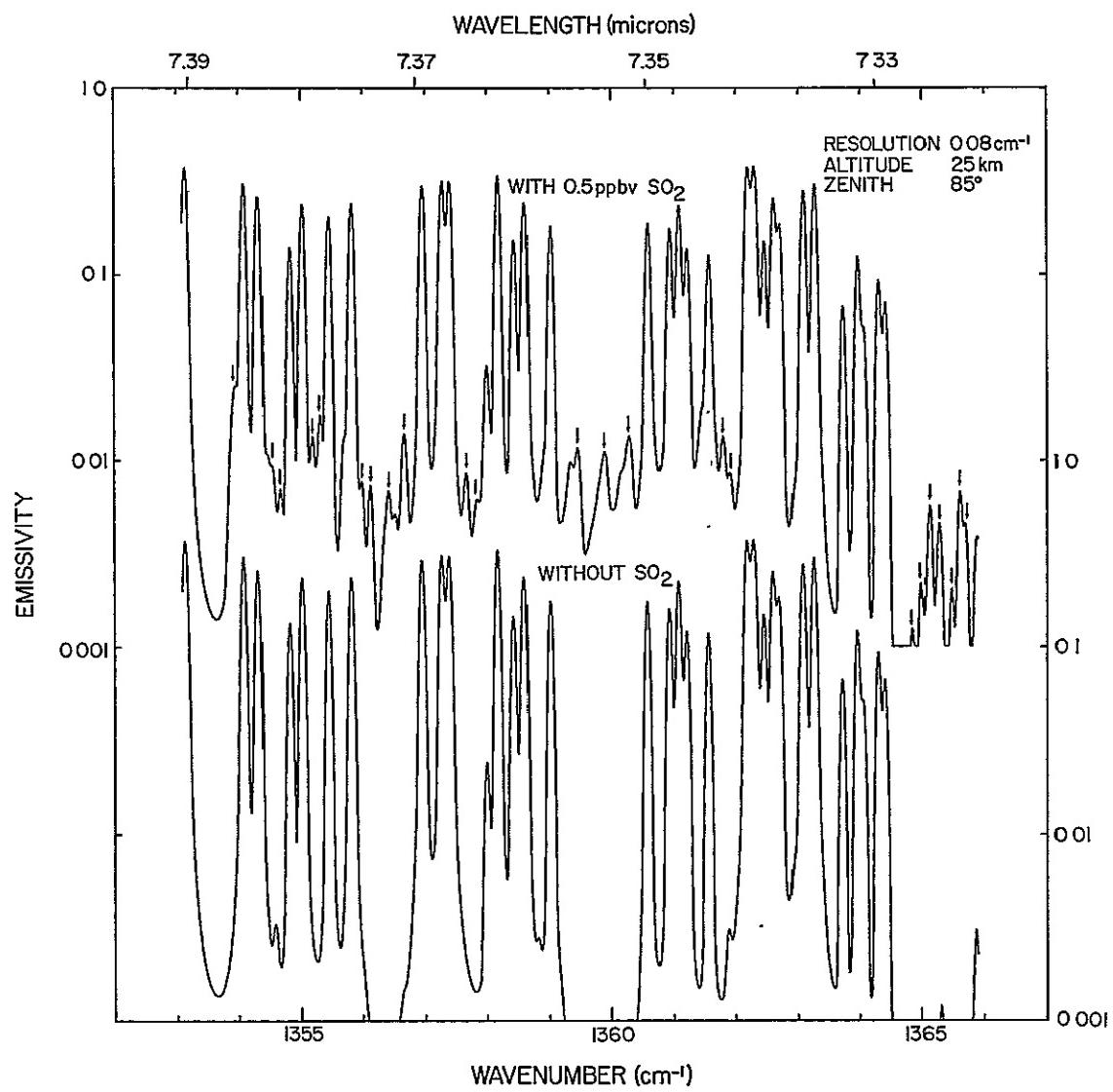


Figure 33.

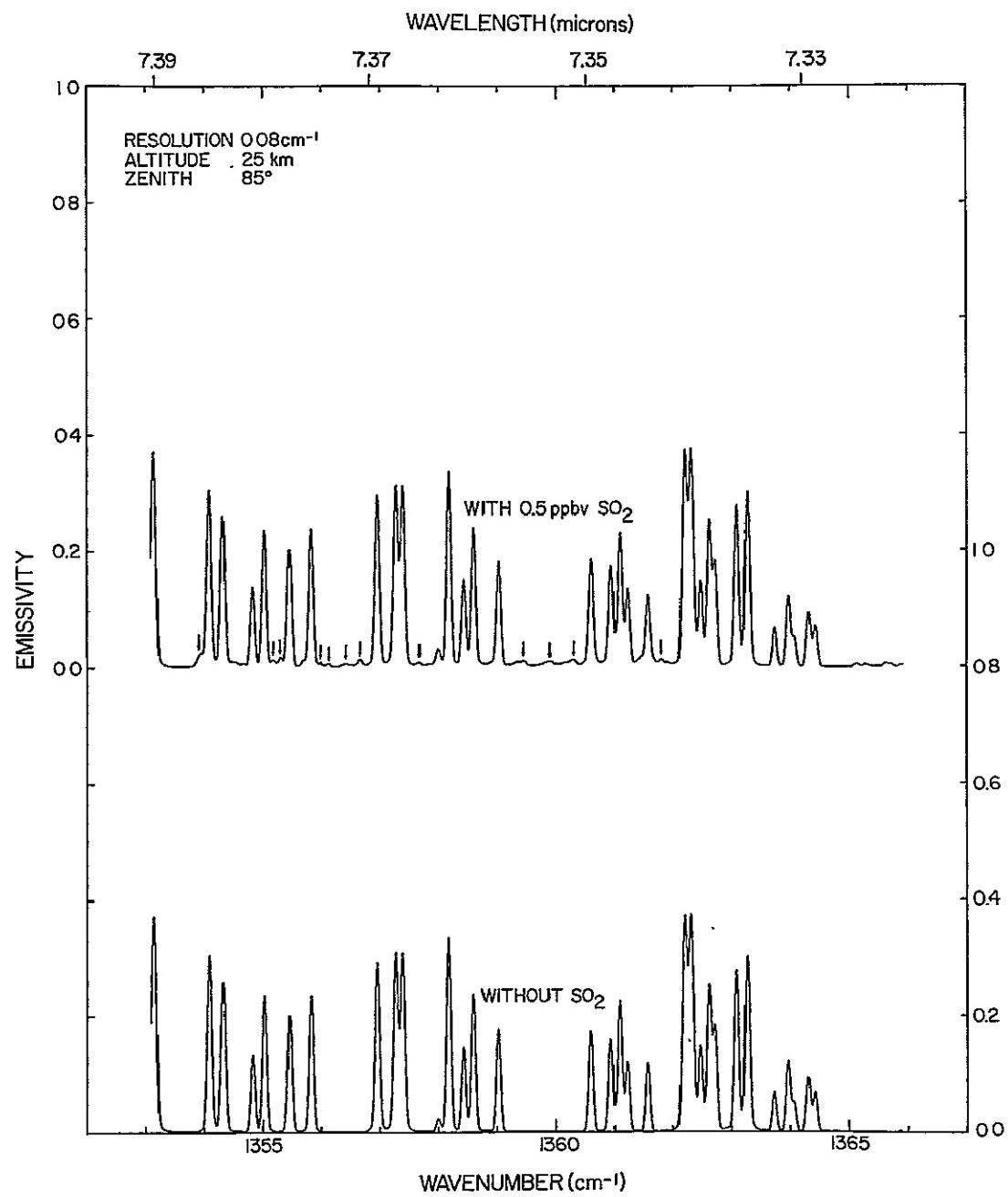


Figure 34.

IV. Discussion

The limits of detectability of the particular constituents discussed above are somewhat arbitrary and are based on the authors' extensive experience with infrared transmittance and radiance measurements and, to a lesser extent, their familiarity with other measurements. It is felt that the detection limits quoted are realistic estimates of what can be done with current instrumentation. The use of a detection limit of 1% absorption in the case of atmospheric absorption using the sun as a source is based on the fact that it is close to the limit achieved in the data published to date.

Examination of the calculated spectra indicate that the combination of resolution and long optical path that can be achieved using the sun as a source during sunset (or sunrise) is a particularly sensitive technique for detection of small amounts of a trace constituent. This technique is particularly useful above the tropopause where interference effects are significantly reduced. As the sensitivity of spectral radiometer systems increases, the advantage of atmospheric absorption (solar spectra) technique, over the atmospheric emission technique, diminishes. When other considerations are taken into account (cost, time of day measurements are needed, data reduction, etc.) quite often the emission technique is better suited for the desired measurement. At the lower altitudes where atmospheric background absorptions start to significantly mask the absorptions due to the desired constituent, the long path absorption cell technique often will be suitable for obtaining data when the other techniques are not. Thus, it is not possible to say that one of the three techniques is the best technique for measuring minor constituents.

Several remarks should be made regarding the spectra presented above. The spectral regions which were selected for use in this study were chosen from the strongest vibration-rotation bands for the molecule having the minimum interference in the region from the "background" constituents. These choices were based on the extensive balloon observations of the variation of atmospheric infrared emission and absorption spectra altitude made by our group along with those of other groups.

No claim is made of completeness in examining any of the parameters of the eight constituents discussed here. Each has many spectral regions which might be useful, while only the stronger spectral transitions were considered here. Most are presumed to have variable mixing ratios with altitude and in the event they are strongly layered, there is undoubtedly a more unique experiment and calculations which could be performed. No consideration of aerosols has been made here, which undoubtedly effects low level emissivity measurements.

It must be emphasized that the concept of detectable limits is not uniquely defined. The calculated spectra demonstrate that the detectable limit depends on the physical parameters of the experimental environment as well as the specific experimental technique. The gas amount denoted in each figure was chosen because it exhibits weak but observable spectral features, allowing an estimate of the minimum detectable amount under the specified experimental conditions. The detectable gas amount depends strongly on the interference with other spectral lines and on the uncertainties in the spectral line parameters of both the background atmosphere and the constituent. These can vary significantly from one case to another, according to path, altitude, spectral region and resolution, from completely isolated

spectral lines of a constituent to lines that are completely masked by the background atmosphere.

In several cases shown here the presence of a constituent on the atmospheric spectrum is manifested not by a unique fine structure, but rather in a change of the intensity of patterns that already exist in the background atmospheric spectrum. Meaningful measurements of trace constituents under such conditions require significantly larger minimum detectable amounts to overcome the uncertainty involved in the measurements and in the theoretical interpretations of the atmospheric spectral patterns.

Finally, the spectral plots given here can be used to estimate the possibility of detection of a different amount of the trace constituents under different experimental conditions by interpolation between the various figures. Such interpolation must be done with care since simple linear interpolation is accurate only if the constituent absorption lines show a few percent absorptivity (<10%) at "infinite" resolution. A similar remark applies to changes in the interfering background lines.

Acknowledgments

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